

Researc

Dry inside: progressive unsaturation within leaves with increasing vapour pressure deficit affects estimation of key leaf gas exchange parameters

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Summary

• Climate change not only leads to higher air temperatures but also increases the vapour pressure deficit (VPD) of the air. Understanding the direct effect of VPD on leaf gas exchange is crucial for precise modelling of stomatal functioning.

• We conducted combined leaf gas exchange and online isotope discrimination measurements on four common European tree species across a VPD range of 0.8–3.6 kPa, while maintaining constant temperatures without soil water limitation. In addition to applying the standard assumption of saturated vapour pressure inside leaves (e_i), we inferred e_i from oxygen isotope discrimination of CO₂ and water vapour.

• e_i desaturated progressively with increasing VPD, consistently across species, resulting in an intercellular relative humidity as low as 0.73 ± 0.11 at the highest tested VPD. Assuming saturation of e_i overestimated the extent of reductions in stomatal conductance and CO₂ mole fraction inside leaves in response to increasing VPD compared with calculations that accounted for unsaturation. In addition, a significant decrease in mesophyll conductance with increasing VPD only occurred when the unsaturation of e_i was considered.

• We suggest that the possibility of unsaturated *e*_i should not be overlooked in measurements related to leaf gas exchange and in stomatal models, especially at high VPD.

Introduction

Driven by the ongoing effects of global change, air vapour pressure deficit (VPD) has increased considerably over recent decades on a global scale and is projected to continue rising in the future (IPCC, 2021; Fang *et al.*, 2022). This increase in VPD, coupled with the rise in global temperature, is considered to have adverse impacts on plant functioning and physiological processes related to leaf gas exchange (Fu *et al.*, 2022; Novick *et al.*, 2024). Such impacts are regarded as important reasons for tree mortality and growth decline in forests (Yuan *et al.*, 2019; Trotsiuk *et al.*, 2021; McDowell *et al.*, 2022; Shekhar *et al.*, 2023).

Variations in VPD influence leaf gas exchange by modifying the humidity gradient from the leaf intercellular air spaces to the atmosphere and by directly driving the movements of guard cells (Peak & Mott, 2011; Buckley, 2019). Studies have frequently indicated that leaf net photosynthesis rate (A_n) and transpiration rate (E) exhibit a nonlinear response to increasing VPD, often increasing to reach a maximum before declining (Farquhar, 1978; Franks *et al.*, 1997; Oren *et al.*, 1999; Shirke & Pathre, 2004; Devi & Reddy, 2018). However, an almost equivalent number of studies have instead revealed monotonic responses of A_n and E,

with A_n decreasing but E increasing with increasing VPD (Dai et al., 1992; Maroco et al., 1997; Yong et al., 1997; Urban et al., 2017). This discrepancy is often attributed to the isohydric and anisohydric behaviours of plants: in response to increasing atmospheric water demand, isohydric plants minimise water loss through transpiration by closing their stomata, while anisohydric plants allow large fluctuations in leaf water status (Hochberg et al., 2018). Additionally, the different VPD-response patterns of A_n and E may be partly explained by the response of leaf gas exchange to temperature, which tends to covary with VPD in nature (Duursma et al., 2014). Therefore, disentangling the direct impact of VPD on leaf gas exchange from that of air temperature is crucial for understanding the underlying mechanisms of how variations in VPD and thus atmospheric drought influence plant functioning in a changing environment (Grossiord et al., 2020; Smith et al., 2020; Schönbeck et al., 2022).

Stomata play a key role in regulating leaf gas exchange, implying that a correct estimation of stomatal behaviour is crucial in models of atmosphere–biosphere carbon dioxide (CO₂) and water (H₂O) fluxes. The standard practice for estimating stomatal conductance (g_s), and furthermore the CO₂ mole fraction in the leaf intercellular air spaces (c_i), in leaf gas exchange

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measurements relies on a longstanding assumption that the air in the intercellular air spaces is saturated with water vapour (Gaastra, 1959; Moss & Rawlins, 1963; von Caemmerer & Farquhar, 1981). In other words, under this assumption, the intercellular vapour pressure (e_i) is equal to the saturation vapour pressure (e_s) at a given leaf temperature (T_{leaf}) . One of the reasons that this assumption is made is simply that e_i cannot be measured directly. Only when the intercellular relative humidity (RH) is approximated as $e_i/e_s = 1$ can g_s be calculated from E and the vapour pressure of the air surrounding the leaf (e_a) , assuming that the boundary layer conductance (gb) is known (Gaastra, 1959). With g_s known, c_i can be calculated from A_n , E and the CO₂ mole fraction of the air surrounding the leaf (c_a) , considering the diffusional processes of CO₂ from the atmosphere to the intercellular air spaces (Moss & Rawlins, 1963; von Caemmerer & Farquhar, 1981). The saturation assumption gained support from studies that indirectly showed that e_i stays near saturation across a variety of conditions (Farquhar & Raschke, 1978; Jones & Higgs, 1980; Sharkey et al., 1982; Cernusak et al., 2019).

An increasing number of studies have, however, shown evidence of unsaturation, that is $e_i/e_s < 1$ (Jarvis & Slatyer, 1970; Ward & Bunce, 1986; Egorov & Karpushkin, 1988; Karpushkin, 1994; Canny & Huang, 2006; Holloway-Phillips et al., 2019). This phenomenon is especially notable in a few studies where there was evidence of progressive unsaturation of e_i with increasing VPD (Cernusak et al., 2018; Wong et al., 2022). The approach recently developed by Cernusak et al. (2018) for calculating intercellular RH provides a useful tool to further test the unsaturation hypothesis. It is based on simultaneous assessments of leaf gas exchange and isotope discrimination. In brief, the approach requires measurements of CO2 and H2O fluxes into and out of a leaf gas exchange cuvette and their carbon (C) and oxygen (O) isotope compositions. During the diffusion of CO₂ from the leaf intercellular air spaces into the chloroplast of C₃ plants, the O atoms in CO₂ exchange with those in the leaf water. This is driven by the enzyme carbonic anhydrase and is thought to reach a complete exchange at the chloroplast surface (Gillon & Yakir, 2000). From the measurements, the oxygen isotopic composition (δ^{18} O) of CO₂ at the sites of carbonic anhydrase activity $(\delta^{18}O_{ca})$ can be determined by considering C¹⁸OO fractionation during CO₂ diffusion from the leaf intercellular air spaces to the chloroplast surface (Cernusak et al., 2004). The ei is involved in this calculation, as it contributes to the determination of c_i . Additionally, the $\delta^{18}O$ of CO_2 in equilibrium with water at the evaporative site $(\delta^{18}O_{ce})$ can be independently estimated from the δ^{18} O of transpired water and the Craig–Gordon equation (Craig & Gordon, 1965), which is also dependent on e_i . Because the chloroplast surface is appressed against the cell walls of the intercellular air spaces in C₃ plants, the δ^{18} O of water at the chloroplast surface can be assumed to be very close to that at the evaporative sites (Farquhar et al., 1993; Busch et al., 2013). Thus, an essential assumption of this approach is that $\delta^{18}O_{ca} = \delta^{18}O_{ce}$; with this assumption, e_i can be allowed to vary iteratively to solve for the value that satisfies the condition $\delta^{18}O_{ca} = \delta^{18}O_{ce}$. If unsaturation of e_i occurs as a rule rather than an exception, this implies that the saturation assumption applied for the analysis of leaf gas exchange measurements can cause biased estimations of g_s and c_i , as well as any further inferences (e.g. photosynthetic capacity) based on the two. Despite existing reports of unsaturated e_i values, there are still discrepancies regarding the extent to which unsaturation can occur with increasing VPD and how much the saturation assumption biases estimations of g_s and c_i . Updating our knowledge on this phenomenon is crucial, as it would fundamentally challenge prevailing understandings of water relations and water transport mechanisms (Buckley & Sack, 2019).

Furthermore, mesophyll conductance (gm) is considered an important factor constraining CO2 diffusion inside the leaf (Cousins et al., 2020; Evans, 2021; Stangl et al., 2022). The VPD responses of g_m may be linked to membrane and cell wall permeability, as well as CO₂ transport through aquaporins in response to VPD (Evans, 2021; Márquez & Busch, 2024). By complementing gas exchange measurements with ¹³CO₂ discrimination (Δ^{13} C) assessments, the total mesophyll conductance to CO₂ diffusing from the intercellular spaces into the chloroplast stroma (g_{m13}) can be estimated. In addition, the mesophyll conductance to CO_2 diffusing from the intercellular spaces to the chloroplast surface (gm18) can be estimated and thus differentiated from gm13 through measurements of C18OO discrimination (Δ^{18} O) (Barbour *et al.*, 2016; Holloway-Phillips et al., 2019; Diao et al., 2024). More importantly, calculations of $g_{\rm m}$ enable additional estimations of the CO₂ mole fractions at the chloroplast surface (c_{ca}) and inside chloroplast stroma (c_c) . Only a few studies have involved investigations of the VPD response of $g_{\rm m}$; some indicated that $g_{\rm m}$ declined significantly with increasing VPD (Bongi & Loreto, 1989; Loucos et al., 2017; Holloway-Phillips et al., 2019), while others did not (Warren, 2008; Perez-Martin et al., 2009; Stangl et al., 2019; Wong et al., 2022). It is also unclear how c_{ca} and c_{c} respond to variations in VPD (Márquez et al., 2023). Since the effects of assuming saturation of e_i (or not) on the estimations of g_s and c_i are carried over to the estimations of g_m , c_{ca} and c_c (Cernusak *et al.*, 2019), it is necessary to investigate how the VPD responses of these variables are affected by assuming or not assuming saturation. Studies targeting this knowledge gap will help us gain a better mechanistic understanding of the VPD response of leaf gas exchange.

In this study, we performed combined classical online leaf gas exchange measurements and online isotope discrimination assessments on potted saplings of four common European tree species (Fagus sylvatica, Picea abies, Quercus petraea, and Tilia cordata) under nonlimiting soil water supply. We carried out the measurements under finely controlled humidity and temperature conditions, spanning a VPD gradient of 0.8-3.6 kPa at both 30°C and 35°C (near optimum temperatures for photosynthesis), with two species assessed at each temperature. We aimed to: (1) investigate the direct VPD responses of leaf gas exchange and isotope discrimination at constant temperatures; (2) evaluate the direct effect of VPD on e_i ; and (3) assess the direct VPD responses of g_s , g_m and leaf internal CO2 mole fractions, considering two scenarios: one assuming a saturated e_i over the VPD gradient, and the other accounting for a progressive unsaturation of e_i at moderate to high VPD. Our findings suggest that future estimations of key leaf gas exchange parameters should account for unsaturation

inside the leaf as a real phenomenon, which facilitates more accurate modelling of plant physiological responses to environmental changes.

Materials and Methods

Plant material

Five 2-yr-old saplings of each Fagus sylvatica L., Picea abies (L.) H. Karst., Quercus petraea (Matt.) Liebl., and Tilia cordata Mill. were used in this study. The saplings were planted into 41 pots with soil mixed with commercial slow-release NPK fertilizer (Osmocote Exact Standard 3-4 M; ICL, Suffolk, UK), then transferred to a climate chamber (Bitzer 6HE-35Y; Kälte 3000 AG, Landquart, Switzerland) to induce leaf flushing. The environmental conditions were air temperature of 25°C, RH of 60%, and light intensity of 110 μ mol m⁻² s⁻¹ during a photoperiod of 18 h; during night-time the air temperature was 15°C and RH 50%. One month later, the plants were transferred to a glasshouse and grew under natural light conditions for 3 months. During this period, the average air temperature was 22.8°C, the average RH was 54.1% and the average air VPD was 1.3 kPa. Throughout the experiment, the plants were watered every 2 d, such that the trays beneath the pots always contained liquid water.

Instrumental setup and calibrations

Online stable isotope techniques, characterized by the combined instrumentation of a leaf gas exchange system (all components from Heinz Walz GmbH, Effeltrich, Germany) and CO2 and H₂O isotope analysers (Delta Ray; Thermo Fisher Scientific Inc., Bremen, Germany and L2120-i; Picarro Inc., Santa Clara, CA, USA, respectively), were used to measure leaf gas exchange and isotope discrimination at the same time. The leaf gas exchange system included a portable photosynthesis control unit (GFS-3000), a leaf gas exchange cuvette (3010-GWK1) and a light source (RGBW-L084). The leaf cuvette was temperature and humidity controlled and supplied with a transversal fan mixing the air inside the cuvette, minimizing the leaf boundary layer resistance and ensuring homogeneous distributions of cuvette temperature (T_{cuv}) , humidity and the CO₂ mole fraction inside the cuvette (c_a) . T_{leaf} was measured inside the leaf cuvette using a thermocouple (3010-CA/TCL). A bypass humidity control system (NFRB0101) served as an additional measure to keep the humidity in the cuvette constant at the preselected value.

The gas exchange system was coupled to the CO₂ and H₂O isotope analysers. With an interposed multiport valve (VICI; Valco Instruments Co. Inc., Houston, TX, USA), and the air streams entering or leaving the leaf cuvette were selected alternately for isotope analysis. Carbon isotopic composition ($\delta^{13}C$) in CO₂ was expressed against Vienna Pee Dee Belemnite, and $\delta^{18}O$ in CO₂ and H₂O vapour were expressed against Vienna Standard Mean Ocean Water. The gas exchange system and the VICI valve were connected via an open split to prevent pressure surges caused by valve switching (Diao *et al.*, 2024).

The gas exchange system, the VICI valve and the plant being measured were placed in a climate chamber (Conviron PGR15; Controlled Environments Ltd, Winnipeg, MB, Canada), and the environmental conditions in the chamber were set to match those in the leaf cuvette. The tube connecting the VICI valve to the isotope analysers was heated to 65°C using a heating band (HTS System AG, Hünenberg, Switzerland). These instrumental arrangements prevented potential condensation inside the instrument system and kept the whole plant at similar environmental conditions. More details of the instrumental setup, with schematics and the operation and calibration procedures, can be found in Diao *et al.* (2024).

Measurement protocol

One day before the measurements, plants of a given species were transferred from the glasshouse to the climate chamber. During the measurements, one foliated branch (for F. sylvatica and P. abies) or one entire leaf (for Q. petraea and T. cordata) was placed in the cuvette. Note that the gas exchange system accepts input VPD values derived from T_{leaf} (i.e. leaf-to-air vapour pressure difference, LAVPD) in Pa kPa^{-1} . This unit needs to be corrected for atmospheric pressure (P_{atm}) to be converted to the unit of kPa. The LAVPD was increased from c. 1 to 4 kPa in steps of c. 1 kPa during the measurements. Afterwards, the air VPD was derived using the measured e_a inside the cuvette and e_s was calculated from the Goff-Gratch equation (Goff & Gratch, 1946) based on the T_{cuv} . This resulted in an ecologically meaningful air VPD gradient of 0.8-3.6 kPa in steps of c. 1 kPa (i.e. 0.79, 1.76, 2.69 and 3.63 kPa; Supporting Information Fig. S1). A VPD of 3.6 kPa is very close to the highest historical annual average VPD in the arid climate zone (3.4 kPa, 1981-2020), as derived from climate models (Fang et al., 2022). Likewise, a mean daily daytime VPD c. 3.6 kPa was observed in temperate forest sites under extreme hot droughts (Shekhar et al., 2023). Under this condition, stem cell expansion, and thus tree growth, may no longer be possible (Zweifel et al., 2021; Novick et al., 2024).

For each VPD step, the target $T_{\rm cuv}$ was 30°C for *Q. petraea* and *P. abies* and 35°C for *F. sylvatica* and *T. cordata*. For each species and VPD step, photosynthetically active radiation at the leaf surface was 800 µmol m⁻² s⁻¹ (light saturation) and $c_{\rm a}$ was 400 µmol mol⁻¹. When the gas exchange rates were at steady state (*c.* 40 min), the gas exchange and the isotope ratio data were recorded.

Calculations of intercellular RH from $\delta^{18}O$

The ratio e_i/e_s was resolved from the coupled measurements of Δ^{18} O and leaf gas exchange based on Cernusak *et al.* (2018) and Cernusak *et al.* (2019). In brief, the method requires measurements of δ^{18} O from both H₂O and CO₂. These measurements provide the necessary information to establish the constraint that the δ^{18} O of CO₂ at the sites of carbonic anhydrase activity matches that in equilibrium with evaporative site water, through which e_i/e_s can be resolved iteratively. A detailed description of the equations used to calculate e_i/e_s is provided in Method S1.

Meanwhile, we refer readers to the simplified algebraic flow path provided in Cernusak *et al.* (2018) and the spreadsheet available in Cernusak *et al.* (2019) as useful tools for understanding and implementing the methodology. A list of abbreviations and symbols used in this study is provided in Table 1.

In the iteration loop of e_i/e_s , an additional unknown parameter, g_{m18} , is involved. Because g_{m18} cannot be solved without knowing c_i beforehand and c_i is only estimated later from e_i/e_s (to be described later) without assuming saturation, it is necessary to use assumed g_{m18} values across the VPD steps. In this study, the lowest measured air VPD was maintained at 0.79 kPa, $e_i/e_s = 1$ could be assumed under this condition. Thus, for each tree which was measured at the lowest VPD, we iteratively varied g_{m18} until the value was found at which $\delta^{18}O_{ca} = \delta^{18}O_{ce}$ with a constraint of $e_i/e_s = 1$. The solved g_{m18} values at the lowest VPD were further assigned to the measurements of the respective tree at higher air VPD conditions. Note that after e_i/e_s values were

Table 1 Definitions of symbols used repeatedly in the main text.

A_n Net photosynthesis rate C_a CO_2 mole fraction of the air surrounding the leaf c_c CO_2 mole fraction in the chloroplast stroma c_{ca} CO_2 mole fraction at the sites of carbonic anhydrase activity c_i CO_2 mole fraction in the leaf intercellular air spaces $\Delta^{13}C$ Photosynthetic C isotope discrimination of CO_2 $\Delta^{13}C$ Observed discrimination against ${}^{13}CO_2$ during photosynthesis $\Delta^{18}O$ Observed discrimination against $C^{18}OO$ during photosynthesis $\delta^{18}O_A$ $\delta^{18}O$ of CO_2 at the sites of carbonic anhydrase activity $\delta^{18}O_{ca}$ $\delta^{18}O$ of CO_2 in equilibrium with evaporative site water $\delta^{18}O_e$ $\delta^{18}O$ of water at evaporative sites in leaves $\delta^{18}O_0$ $\delta^{18}O$ of transpired water vapour
c_a CO_2 mole fraction of the air surrounding the leaf c_c CO_2 mole fraction in the chloroplast stroma c_{ca} CO_2 mole fraction at the sites of carbonic anhydrase activity c_i CO_2 mole fraction in the leaf intercellular air spaces $\Delta^{13}C$ Photosynthetic C isotope discrimination of CO_2 $\Delta^{13}C$ Observed discrimination against ${}^{13}CO_2$ during photosynthesis $\Delta^{18}O_{obs}$ Observed discrimination against $C^{18}OO$ during photosynthesis $\Delta^{18}O_{abc}$ $\delta^{18}O$ of CO_2 taken up by net photosynthesis $\delta^{18}O_{ca}$ $\delta^{18}O$ of CO_2 in the sites of carbonic anhydrase activity $\delta^{18}O_{ca}$ $\delta^{18}O$ of CO_2 in equilibrium with evaporative site water $\delta^{18}O_{c}$ $\delta^{18}O$ of transpired water vaporative
c_c CO2 mole fraction in the chloroplast stroma c_{ca} CO2 mole fraction at the sites of carbonic anhydrase activity c_i CO2 mole fraction in the leaf intercellular air spaces $\Delta^{13}C$ Photosynthetic C isotope discrimination of CO2 $\Delta^{13}C_{obs}$ Observed discrimination against ${}^{13}CO_2$ during photosynthesis $\Delta^{18}O_{obs}$ Observed discrimination against $C^{18}OO$ during photosynthesis $\Delta^{18}O_{obs}$ Observed discrimination against $C^{18}OO$ during photosynthesis $\delta^{18}O_{ca}$ $\delta^{18}O$ of CO2 taken up by net photosynthesis $\delta^{18}O_{ca}$ $\delta^{18}O$ of CO2 in equilibrium with evaporative site water $\delta^{18}O_{ca}$ $\delta^{18}O$ of water at evaporative sites in leaves $\delta^{18}O_{ca}$ $\delta^{18}O$ of transpired water vapour
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$ \Delta^{13}C Photosynthetic C isotope discrimination of CO_2 \Delta^{13}C_{obs} Observed discrimination against 13CO_2 during photosynthesis \Delta^{18}O Photosynthetic O isotope discrimination of CO_2 \Delta^{18}O_{obs} Observed discrimination against C18OO during photosynthesis \delta^{18}O_A \delta^{18}O of CO_2 taken up by net photosynthesis \delta^{18}O_{ca} \delta^{18}O of CO_2 at the sites of carbonic anhydrase activity \delta^{18}O_{ce} \delta^{18}O of CO_2 in equilibrium with evaporative site water \delta^{18}O_r \delta^{18}O of transpired water vapour$
$ \Delta^{13}C_{obs} $ Observed discrimination against ${}^{13}CO_2$ during photosynthesis $ \Delta^{18}O $ Photosynthetic O isotope discrimination of CO_2 $ \Delta^{18}O_{obs} $ Observed discrimination against $C^{18}OO$ during photosynthesis $ \delta^{18}O_A $ $\delta^{18}O$ of CO_2 taken up by net photosynthesis $ \delta^{18}O_{ca} $ $\delta^{18}O$ of CO_2 at the sites of carbonic anhydrase activity $ \delta^{18}O_{ce} $ $\delta^{18}O$ of CO_2 in equilibrium with evaporative site water $ \delta^{18}O_r $ $\delta^{18}O$ of yater at evaporative sites in leaves $ \delta^{18}O_r $ $\delta^{18}O$ of transpired water vapour
$ \Delta^{18}O \\ \Delta^{$
$\begin{array}{lll} \Delta^{18} O_{obs} & Observed discrimination against C^{18} OO during photosynthesis \\ \delta^{18} O_A & \delta^{18} O \ of \ CO_2 \ taken \ up \ by \ net \ photosynthesis \\ \delta^{18} O_{ca} & \delta^{18} O \ of \ CO_2 \ at \ the \ sites \ of \ carbonic \ anhydrase \ activity \\ \delta^{18} O_{ce} & \delta^{18} O \ of \ CO_2 \ in \ equilibrium \ with \ evaporative \ site \ water \\ \delta^{18} O_e & \delta^{18} O \ of \ water \ at \ evaporative \ sites \ in \ leaves \\ \delta^{18} O_r & \delta^{18} O \ of \ transpired \ water \ vapour \end{array}$
$\delta^{18}O_A$ $\delta^{18}O$ of CO_2 taken up by net photosynthesis $\delta^{18}O_{ca}$ $\delta^{18}O$ of CO_2 at the sites of carbonic anhydrase activity $\delta^{18}O_{ce}$ $\delta^{18}O$ of CO_2 in equilibrium with evaporative site water $\delta^{18}O_e$ $\delta^{18}O$ of water at evaporative sites in leaves $\delta^{18}O_r$ $\delta^{18}O$ of transpired water vapour
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$\delta^{18}O_{ce}$ $\delta^{18}O$ of CO_2 in equilibrium with evaporative site water $\delta^{18}O_e$ $\delta^{18}O$ of water at evaporative sites in leaves $\delta^{18}O_r$ $\delta^{18}O$ of transpired water vapour
$\delta^{18}O_e$ $\delta^{18}O$ of transpired water vapour
$\lambda^{(0)}$ O_{r} $\lambda^{(0)}$ O of transpired water vapour
$\delta^{10}O_i = \delta^{10}O$ of CO_2 in the intercellular air spaces
E Transpiration rate
e _a Vapour pressure of the last intercellular sizeness
ei Vapour pressure in the leaf intercentuar air spaces
temperature
g _b Leaf boundary layer conductance to water vapour
$g_{\rm m}$ Mesophyll conductance to $\rm CO_2$
g_{m13} Mesophyll conductance to CO ₂ estimated from ¹³ C
measurements, representing the mesophyll conductance from
leaf intercellular air spaces to sites of carboxylation (i.e. total
mesophyll conductance)
g_{m18} Mesophyli conductance to CO_2 estimated from ^{12}O
losf intercellular air spaces to sites of earbonic anbudrase (i.e.
sell wall and plasma membrane conductance)
a Stomatal conductance to water vapour
gs Stolliatal colluctance to water vapour
to the atmosphere
g_{tc} Conductance to CO ₂ from leaf intercellular air spaces to the
aumosphere PH Polativo humidity
The Curvette temperature
$T_{\rm cuv}$ Cuvelle lemperature
VPD Air vanour pressure deficit

Symbols that appear only alongside their definition in the text are not included here.

estimated, e_i was further derived from e_s at T_{leaf} according to the temperature– e_s relationship of Goff & Gratch (1946). The data used in this study are provided in the Dataset S1.

Calculations of stomatal conductance and intercellular $\rm CO_2$ mole fraction with and without assumed saturation

 $A_{\rm n}$ and *E* values were calculated by the Walz GFS-3000 gas exchange system, according to von Caemmerer & Farquhar (1981). For each individual measurement, $g_{\rm s}$ and $c_{\rm i}$ were calculated using an identical set of equations and either substituting $e_{\rm i}$ with $e_{\rm s}$ ($e_{\rm i}/e_{\rm s} = 1$) to refer to the saturation assumption at VPD of 0.8 kPa or using $e_{\rm i}$ independently derived from the Δ^{18} O measurements to account for unsaturation at VPD > 0.8 kPa as follows (von Caemmerer & Farquhar, 1981).

The variable g_s was given as:

$$g_s = \frac{1}{\frac{1}{g_t} - \frac{1}{g_b}}$$
 Eqn (

where g_t is the total conductance to water vapour from leaf intercellular air spaces to the atmosphere, and g_b is assumed to be 3.5 mol H₂O m⁻² s⁻¹ for the chamber and fan speed used in our measurements (Cernusak *et al.*, 2019; Diao *et al.*, 2024).

The variable g_t was given as:

$$g_{t} = \frac{E \times \left(1 - \frac{e_{i} + e_{a}}{2P_{atm}}\right)}{\frac{e_{i} - e_{a}}{P_{atm}}}$$
Eqn 2

The variable c_i was given as:

$$c_{\rm i} = \frac{\left(g_{\rm tc} - \frac{E}{2}\right) \times c_{\rm a} - A_{\rm n}}{g_{\rm tc} + \frac{E}{2}}$$
 Eqn 3

where g_{tc} is the total conductance to CO₂, expressed as $1/(1.6/g_s + 1.37/g_b)$.

Calculations of chloroplastic CO_2 mole fraction and mesophyll conductance

The CO₂ mole fraction at the chloroplast surface (c_{ca}), the assumed site of carbonic anhydrase activity, can be derived from δ^{18} O measurements according to Barbour *et al.* (2016).

$$c_{\rm ca} = c_{\rm i} \left(\frac{\delta^{18} O_{\rm i} - \overline{a} - \alpha_{\rm ac} \delta^{18} O_{\rm A}}{\delta^{18} O_{\rm ce} - \overline{a} - \alpha_{\rm ac} \delta^{18} O_{\rm A}} \right)$$
 Eqn 4

where α_{ac} is the ¹⁸O/¹⁶O fractionation factor for CO₂ diffusing across the boundary layer and stomata, defined as $1+\bar{a}$. The variable \bar{a} was calculated as described in Diao *et al.* (2024). Here, c_i was derived by employing the saturation assumption across the VPD steps; Eqn 4 was, therefore, used to calculate c_{ca} assuming saturation of c_i .

In the calculations associated with unsaturation, c_i cannot be derived from the iteration if c_{ca} is not known first. Therefore, the c_{ca} which accounted for unsaturation of e_i was calculated (Eqn 5) by

assuming that the g_{m18} values determined at the lowest VPD step for each leaf remained unchanged at the other VPD steps and by using the c_i for which unsaturation was taken into account (Eqn 3).

$$c_{\rm ca} = c_{\rm i} - \frac{A_{\rm n}}{g_{\rm m18}}$$
 Eqn 5

The variable g_{m13} was expressed as:

$$g_{m13} = \frac{1+t}{1-t} \left(\frac{b-a_m - \frac{a_b}{a_e a_R} e^{t} \frac{\mathcal{R}_{day}}{A_n}}{\Delta^{13} C_i - \Delta^{13} C_{obs}} \right) \frac{A_n}{c_a}$$
Eqn 6

where $a_{\rm m}$ is the ¹³C/¹²C fractionation for CO₂ dissolution and diffusion from the intercellular air spaces to the sites of carboxylation in the chloroplasts (1.8‰), and b is the ${}^{13}C/{}^{12}C$ fractionation associated with carboxylation (30%). $e' = e + e^*$, where e is the ${}^{13}C/{}^{12}C$ fractionation during day respiration (-3‰) and e^* is an additional apparent fractionation accounting for the difference in substrate composition between freshly assimilated C and old C pools (Wingate et al., 2007). e* was assumed to be zero in our case (Diao *et al.*, 2024). $\alpha_{\rm b}$ and $\alpha_{\rm e}$ are defined as 1 + b and 1 + e, respectively. R_{dav} is the day respiration rate and assumed to be half of the dark respiration rate (R_{dark}) (Adnew *et al.*, 2020). The R_{dark} measurements were described in Diao *et al.* (2024). $\alpha_{\rm R} = 1 + (R_{\rm dav}/A_{\rm n})(e/\alpha_{\rm e})$ (Busch *et al.*, 2020). The ternary correction term, t, was calculated according to Farquhar & Cernusak (2012). $\Delta^{13}C_{obs}$ is the observed discrimination against ^{13}C during CO2 assimilation, which was calculated following Evans et al. (1986). $\Delta^{13}C_i$ is the modelled C isotope discrimination, which was calculated according to Busch et al. (2020). Detailed descriptions of the calculations are presented in Diao et al. (2024). Note that the c_i values calculated both with the saturation assumption and with the unsaturation of e_i accounted for were used in the calculations of g_{m13} for comparison.

The CO₂ mole fraction inside the chloroplast stroma (c_c) was then calculated separately for the saturation and unsaturation cases, as:

$$c_{\rm c} = c_{\rm i} - \frac{A_{\rm n}}{g_{\rm m13}}$$
 Eqn 7

Statistical analysis

All statistical analyses were conducted using R software v.4.3.1 (R Core Team, 2023). The responses of the studied variables to air VPD were analysed using linear mixed models with the IMERT-EST package in R (Kuznetsova *et al.*, 2017), with air VPD as a fixed effect and both species and T_{cuv} groups (i.e. $T_{cuv} = 30^{\circ}$ C and 35°C) as random effects. Because the T_{cuv} groups were treated as random effects in our statistical analyses, and we did not observe apparent separation of the VPD-response curves based on T_{cuv} in our dataset, we confidently conclude that temperature-specific patterns were not evident in this study and that the observed physiological responses were purely driven by VPD. The between-group sum of squares of an analysis of variance (ANOVA) was used to compare the differences in interspecies variabilities between g_s calculated with and without assuming saturation of e_i . Two-sample *t*-tests were performed to test the significance of differences in g_s and g_{m13} when assuming saturation of e_i and when accounting for unsaturation of e_i . For $\Delta^{13}C_{obs}$ and the observed discrimination against ¹⁸O during CO₂ assimilation ($\Delta^{18}O_{obs}$), a filter ($\xi < 30$, defined in Eqn S6) was applied to exclude data with low confidence in the isotopic discrimination estimates (Diao *et al.*, 2024).

Results

Experimental conditions

 $T_{\rm cuv}$ was kept constant along the VPD gradient, with a mean value (±1 SD) of $30.0 \pm 0.0^{\circ}$ C for *P. abies* and *Q. petraea*, and of $35.0 \pm 0.0^{\circ}$ C for *F. sylvatica* and *T. cordata* (Fig. S1). $T_{\rm leaf}$ was slightly higher than $T_{\rm cuv}$, with a mean value of $30.3 \pm 0.1^{\circ}$ C for *P. abies* and *Q. petraea*, and of $35.2 \pm 0.1^{\circ}$ C for *F. sylvatica* and *T. cordata*. For these two groups of $T_{\rm cuv}$, VPD was increased stepwise from 0.8 ± 0.1 to 3.6 ± 0.1 kPa, resulting in decreases in RH from $78.3\% \pm 2.8\%$ to $13.4\% \pm 2.0\%$ and from $78.5\% \pm 1.1\%$ to $32.2\% \pm 1.3\%$, respectively. These relationships match well with the theoretical relationship between VPD and RH at constant temperatures (Fig. S1), suggesting precise temperature and humidity control in the gas exchange system.

Responses of A_n , E, and C and O isotope discriminations to increasing VPD

The directions of responses of A_n and E to increasing air VPD were consistent among the four studied species (Fig. 1). Based on the analyses of linear mixed models and averaged across the species, A_n decreased significantly from 5.4 ± 1.6 to $3.9 \pm 1.1 \,\mu$ mol CO₂ m⁻² s⁻¹ with increasing air VPD (Fig. 1a; F=17.51, P<0.001), whereas E increased significantly from 1.2 ± 0.4 to $1.8 \pm 0.8 \,\mathrm{mmol} \,\mathrm{H_2O} \,\mathrm{m^{-2} \, s^{-1}}$ with increasing air VPD (Fig. 1b; F=16.19, P<0.001). Note that the VPD response of A_n was much less pronounced for F. sylvatica than for the other three species, with a slope of $-0.22 \,\mu$ mol CO₂ m⁻² s⁻¹ kPa⁻¹ for F. sylvatica and a slope of $-0.71 \,\mu$ mol CO₂ m⁻² s⁻¹ kPa⁻¹ averaged across the remaining species (Fig. 1a).

We found significant responses of $\Delta^{13}C_{obs}$ (F=7.81, P=0.007) and $\Delta^{18}O_{obs}$ (F=23.73, P<0.001) to increasing air VPD across the species (Fig. S2). On average, $\Delta^{13}C_{obs}$ decreased with increasing air VPD, from 19.7 ± 1.9 to $17.3 \pm 2.1\%$ (Fig. S2a), while $\Delta^{18}O_{obs}$ increased from 35.0 ± 10.4 to $45.9 \pm 11.3\%$ (Fig. S2b). However, for *F. sylvatica*, the linear relationship between $\Delta^{13}C_{obs}$ and air VPD was not statistically significant (Fig. S2a; P=0.98).

Responses of RH inside the leaf to VPD

We constrained the values of the leaf intercellular RH as $e_i/e_s = 1$ at the lowest measured VPD (< 1 kPa), ensuring that these values explicitly align with the assumption that the air inside the leaf is



saturated with water vapour under low VPD (Fig. 2). At VPD > 1 kPa, intercellular RH was calculated from the δ^{18} O measurements (see Materials and Methods section). This approach takes unsaturation of e_i into account. Across the species, intercellular RH decreased significantly with increasing air VPD (F=91.83, P < 0.001), suggesting a progressive unsaturation inside the leaf. Specifically, across the species which were measured at 30°C and 35°C, a linear regression showed that with every 1 kPa increase in air VPD, intercellular RH decreased by 0.10 ± 0.01 (slope \pm SE for the regression model). At the highest measured VPD (i.e. 3.6 ± 0.1 kPa), the intercellular RH was 0.73 ± 0.11 .

VPD responses of $g_{\rm s}$ and $g_{\rm m}$ with and without assuming saturation of $e_{\rm i}$

The g_s showed a consistent pattern of decrease with increasing VPD, both with and without the saturation assumption (Fig. 3). Across all species, a reciprocal function fit the responses well (P < 0.001 and P = 0.01 under saturation assumed and unsaturation accounted for, respectively), and the decreases were statistically significant when both species and T_{cuv} groups were included in the models as random effects (F = 91.92 and 9.02, respectively; P < 0.01 for both). However, the VPD-response pattern of the g_s calculated with the saturation assumption (Fig. 3a) was more distinct than that calculated accounting for unsaturation (Fig. 3b). The difference in average gs between the lowest and highest VPD was much smaller (0.05 mol $H_2O m^{-2} s^{-1}$) and the inter-species variability of the values was much higher (0.06; between-species sum of squares, ANOVA) for the case accounting for unsaturation (Fig. 3b), than when saturation was assumed (0.09 mol H_2O $m^{-2} s^{-1}$, 0.03; Fig. 3a). Moreover, for VPD > 0.8 kPa, the g_s calculated accounting for unsaturation (Fig. 3b) was significantly higher (0.03 mol H₂O m⁻² s⁻¹; P < 0.001, *t*-test) than the gs when saturation was assumed (Fig. 3a).

No significant response of g_{m13} to increasing air VPD was found when saturation of e_i was assumed, although the smoothed trend indicated somewhat higher g_{m13} values on average when VPD was also high (Fig. 4a). By contrast, when unsaturation was accounted for, g_{m13} decreased significantly with increasing VPD across the species (F = 6.95, P = 0.01), ranging from 0.10 ± 0.06 mol CO₂

Fig. 1 Responses of (a) net photosynthesis rate (A_n) and (b) transpiration rate (*E*) to increasing air vapour pressure deficit (VPD) in four tree species. Each dashed line represents the linear trend for one tree species. The solid line and the grey band represent the linear trend averaged across all tree species and its 95% confidence interval, respectively. Different symbols and colours indicate different tree species (n = 5 per VPD step).



Fig. 2 Response of leaf intercellular relative humidity (e_i/e_s) to increasing air vapour pressure deficit (VPD) in four tree species. The e_i/e_s values were estimated from measurements of oxygen isotope discrimination in CO₂ and transpired water vapour. At the lowest VPD step, $e_i/e_s = 1$ was applied to meet the assumption of vapour saturation inside the leaf. Each dashed line represents the linear trend for one tree species. The solid line and the grey band represent the linear trend averaged across all tree species and its 95% confidence interval, respectively. Different symbols and colours indicate different tree species (n = 5 per VPD step).

 $m^{-2}s^{-1}$ at the lowest measured VPD to 0.06 ± 0.03 mol CO₂ $m^{-2}s^{-1}$ at the highest measured VPD (Fig. 4b). However, for *F. sylvatica*, the linear relationship between g_{m13} and air VPD was not statistically significant (Fig. 4b; P = 0.90).

Responses of the CO_2 mole fraction inside the leaf to VPD with and without assuming saturation of e_i

Across the air VPD gradient, the average CO_2 mole fractions within leaves decreased progressively from the outside

Fig. 3 Responses of stomatal conductance (g_s) to increasing air vapour pressure deficit (VPD) in four tree species. Saturation of the leaf intercellular vapour pressure (e_i) was assumed in calculations for the g_s values in (a), whereas unsaturation was taken into account in calculations for the g_s values in (b). Each dashed line represents the regression of a reciprocal function (y = 1/x) for one species. Different symbols and colours indicate different tree species (n = 5 per VPD step). The solid line and the grey band represent the regression of a reciprocal function (y = 1/x) averaged across all tree species and its 95% confidence interval, respectively.





Fig. 4 Responses of mesophyll conductance to CO_2 estimated from ¹³C measurements (g_{m13}) to increasing air vapour pressure deficit (VPD) in four tree species. Saturation of the leaf intercellular vapour pressure (e_i) was assumed in calculations for the g_{m13} values in (a), whereas unsaturation was taken into account in calculations for the g_{m13} values in (b). In (a), each dashed line represents a local polynomial regression fitting for one tree species. The solid line and the grey band represent a local polynomial regression fitting averaged across all species and its 95% confidence interval, respectively. In (b), each dashed line and the grey band represents the linear trend for one species. The solid line and the grey band represent the linear trend averaged across all species and its 95% confidence interval, respectively. In both panels, different symbols and colours indicate different tree species (n = 5 per VPD step).

in: $c_i/c_a > c_{ca}/c_a > c_c/c_a$ (Fig. 5). This pattern held both with and without the saturation assumption (both P < 0.001). With the saturation assumption, c_i/c_a (F = 137.07, P < 0.001), c_{ca}/c_a (F = 11.51, P = 0.001), and c_c/c_a (F = 24.02, P < 0.001) decreased significantly with increasing air VPD (Fig. 5a-c), but the decreasing trend in c_{ca}/c_a (Fig. 5b) was not as substantial as that of c_i/c_a and c_c/c_a (Fig. 5a, c). Accounting for unsaturation, c_i/c_a (F = 10.67, P = 0.002) and c_c/c_a (F = 14.47, P < 0.001) decreased significantly with the increase in air VPD, but this was not found for c_{ca}/c_a (F = 3.46, P = 0.07) (Fig. 5d-f).

Regarding the rates of changes in the CO₂ mole fractions with increasing air VPD, c_i/c_a decreased less with increasing VPD when unsaturation was taken into account than when saturation was assumed, as shown by the different slopes of the linear fittings (-0.02, Fig. 5d vs -0.08, Fig. 5a). For c_{ca}/c_a and c_c/c_a , the

decreasing rates were not distinctly different when saturation was assumed (Fig. 5b,c) or not (Fig. 5e,f).

Discussion

Progressive unsaturation of e_i with increasing VPD

We quantitatively estimated e_i/e_s values of the four studied tree species as they were exposed to increasing VPD, using measurements of δ^{18} O of CO₂ and water vapour (Cernusak *et al.*, 2018). We found strong signals of unsaturation and the degree of unsaturation increased linearly with the increase in VPD (Fig. 2). In agreement with the evidence of unsaturation shown in this study, Cernusak *et al.* (2018) reported that intercellular RH declined to values in the range of 0.9 in *Pinus edulis* at *c.* 5 kPa VPD and 0.8 in *Juniperus monosperma* at *c.*



Fig. 5 Air vapour pressure deficit (VPD) responses of the ratio of intercellular to ambient CO₂ mole fraction (c_i/c_a ; a, d); the ratio of chloroplast surface to ambient CO₂ mole fraction (c_i/c_a ; b, e); and the ratio of chloroplast stroma to ambient CO₂ mole fraction (c_c/c_a ; c, f). Unsaturation of the leaf intercellular vapour pressure (e_i) was taken into account in calculations for (d–f) but not for (a–c). Each dashed line represents the linear trend for one species. Different symbols and colours indicate different tree species (n = 5 per VPD step). The solid line and the grey band represent the linear trend averaged across all species and its 95% confidence interval, respectively.

4 kPa VPD; Holloway-Phillips et al. (2019) reported that intercellular RH declined to c. 0.95 in Vicia faba at c. 2 kPa VPD; and Wong et al. (2022) reported that intercellular RH decreased to 0.8 in Gossypium hirsutum and to 0.9 in Helianthus annuus at c. 2 kPa VPD. The extent of the decrease in intercellular RH as VPD rises in our study (i.e. decrease of 0.1 per 1 kPa increase in VPD) shows close agreement with the values reported by Holloway-Phillips et al. (2019) and Wong et al. (2022). Even though this decline is more severe than what was reported by Cernusak et al. (2018), it is still lower than the slope of 0.23 kPa^{-1} found in the transgenic Populus × canescens transformed with the Arabidopsis thaliana mutant *abi1* gene, which has a defective stomatal closure response to VPD, therefore, reflecting an extreme case of unsaturation (Cernusak et al., 2019). In addition, the results of the studies mentioned above and this study together suggest that the saturation assumption is generally reliable for leaf gas exchange measurements at VPD < 1 kPa. Still, further studies on direct observations of stomata are required to assess the

degree of unsaturation across a range of VPD in different species acclimatised to diverse environmental conditions.

The reliability of the method we used to estimate e was critically evaluated in Cernusak et al. (2018). They concluded that a decline in e_i below saturation as VPD increased is the most likely explanation for the progressively higher $\delta^{18}O_i$ values compared with $\delta^{18}O_{ce}$, suggesting the robustness of the method in capturing potential unsaturation of e_i . In addition, as acknowledged in Cernusak *et al.* (2018), the method we used to estimate e_i is sensitive to the choice of gm18 values. Recently, Holloway-Phillips et al. (2019) developed a two-source δ^{18} O method, which does not need assumed g_{m18} values to estimate e_i ; they found a decrease in g_{m18} with increasing air VPD (g_{m18} = -0.13VPD + 0.97). We further performed a sensitivity analysis with g_{m18} values that decrease with increasing VPD, based on the observations in Holloway-Phillips et al. (2019) (see Supporting Information). We found that the e_i calculated with VPD-sensitive gm18 values generally decreased more strongly with increasing VPD than in the case of constant g_{m18} (Fig. S3).

Moreover, an essential assumption of the oxygen isotope method is that $\delta^{18}O_{ca} = \delta^{18}O_{ce}$, which implies that carbonic anhydrase activity is sufficiently high to drive full equilibration between CO₂ and water at the evaporative sites across the VPD gradient. By contrast, the c_i difference technique developed by Wong *et al.* (2022) does not require this assumption. The comparison, illustrated in Extended Data fig. 2 of Wong *et al.* (2022), shows that both methods produced consistent decreases in e_i/e_s with increasing VPD. Thus, our observation that e_i/e_s progressively declines below unity as air VPD increases, as shown in Fig. 2, remains valid.

Differences in the VPD responses of g_s and g_m with and without assuming saturation of e_i

 $A_{\rm n}$ and E are directly measured fluxes and thus cannot be affected by whether saturation or unsaturation is assumed. A_n decreased but E increased with increasing VPD (Fig. 1), both findings that are consistent with previously reported VPD-response patterns conducted at constant temperatures (Dai et al., 1992; Maroco et al., 1997; Yong et al., 1997; Urban et al., 2017). Independent of whether saturation of e_i was assumed, g_s decreased as VPD increased, with a sharper drop in the lower VPD range (Fig. 3). This VPD-response pattern of g_s is consistent with previous findings (Dai et al., 1992; Maroco et al., 1997; Yong et al., 1997; Cernusak et al., 2019). However, the saturation assumption resulted in a considerable underestimation of g_s, and after accounting for unsaturation, the decrease in gs with increasing VPD became less evident (Fig. 3). These results imply that the impact strength of VPD to induce stomatal closure may have been overestimated in previous studies, which generally assumed saturation of e_i . As a result of the less pronounced VPD response of g_s when unsaturation was accounted for, the intrinsic water use efficiency (i.e. $iWUE = A_n/g_s$) generally became unresponsive to variations in VPD after accounting for the unsaturation of e_i (Fig. <u>\$4</u>).

Both the vapour pressure gradient from the leaf to the air and g_t determine the VPD response of E (Fig. 1b). For the former, e_i/e_s plays a role; the latter represents the influence of g_s , considering that g_b was held constant in our experimental conditions. Following Cernusak *et al.* (2024), the extent to which variations in E as air VPD increased were driven by e_i/e_s and g_s are shown in Fig. S5. For the two species measured at a T_{cuv} of 30°C (*P. abies* and *Q. petraea*), the reduction in potential *E* reached *c.* –25% due to reduced e_i/e_s and *c.* –46% due to reduced g_t . For the two species measured at a T_{cuv} of 35°C (*F. sylvatica* and *T. cordata*), the reduction in potential *E* reached *c.* –45% due to reduced e_i/e_s and *c.* –29% due to reduced g_t . These results suggest that both e_i/e_s and g_s make important contributions to driving *E* in response to increasing air VPD and that their relative importance could depend on temperature.

Accounting for unsaturation had a greater impact on the VPD-response pattern of g_{m13} than on that of g_s . The variable g_{m13} exhibited a shift from no significant response to VPD when e_i was assumed to be saturated (Fig. 4a) to a significant decreasing trend in response to increasing VPD when unsaturation of e_i was

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considered (Fig. 4b). The result for g_{m13} under unsaturation is in line with findings from some previous studies (Bongi & Loreto, 1989; Loucos *et al.*, 2017; Holloway-Phillips *et al.*, 2019). The decrease in g_m with increasing VPD could be related to decreases in membrane and cell wall permeability and to CO₂ transport by aquaporins (Evans, 2021; Márquez & Busch, 2024).

Differences in the VPD responses of the CO_2 mole fraction inside the leaf with and without assuming saturation of e_i

We found significant decreases in c_1/c_2 and c_2/c_2 with increasing VPD, both with and without the saturation assumption (Fig. 5). It could be that the combined action of reduced g_s and g_m (Figs 3, 4b) limited the CO₂ supply inside chloroplasts, thus leading to suppressed photosynthesis (Fig. 1a). $\Delta^{13}C_{obs}$ and $\Delta^{18}O_{obs}$ (Fig. S2) were unaffected by whether saturation is assumed, as they were directly derived from the isotope measurements. The decrease in $\Delta^{13}C_{obs}$ with increasing VPD (Fig. S2a) independently supports the results of c_c/c_a , as a decrease in c_c results in decreased $\Delta^{13}C_{obs}$ (Farquhar *et al.*, 1982; Evans *et al.*, 1986). While c_{ca} taking into account unsaturation was calculated from assumed constant g_{m18} values (Eqn 5), c_{ca} assuming saturation was not (Eqn 4). Nevertheless, in both cases, c_{ca}/c_a only slightly decreased with increasing air VPD (Fig. 5b,e), differing from the results for c_i/c_a and c_c/c_a (Fig. 5). The maintained c_{ca}/c_a at high VPD is supported by the increase in $\Delta^{18}O_{obs}$ with increasing VPD (Fig. S2b), indicating continuous CO₂ diffusion into leaves which exchanged O atoms at the sites of carbonic anhydrase activity with leaf water and progressively isotopically enriched as VPD increased (Gillon & Yakir, 2000).

We found that taking unsaturation into account induced a great change in the VPD-response pattern of c_1/c_2 (Fig. 5a,d). The saturation assumption leads to an overestimation of the role of VPD in inducing stomatal closure and, consequently, in reducing c_i , as shown here. This would cause biased estimations of other parameters that involve c_i, for example, maximum rates of carboxylation and electron transport, as well as the rate of respiration in the presence of light derived from $A_n - c_i$ curves. While corrections of c_i based on the unsaturation conditions of leaf intercellular spaces are not yet feasible in many cases, care should be taken regarding data interpretation involving c_i derived from conventional leaf gas exchange measurements. Surprisingly, even though a dramatic difference was shown for the estimations of g_{m13} with and without assuming saturation (Fig. 4), the VPD responses of c_c were not so different under the two scenarios (Fig. 5c,f). This result questions the significance of g_m in constraining the diffusion of CO₂ inside leaves, which is in agreement with the rising awareness that $g_{\rm m}$ and c_c do not necessarily covary (Bahar et al., 2018; Diao et al., 2024).

Implications for leaf water potential and water transport mechanisms

The unsaturation of e_i implies very negative water potentials in the intercellular air spaces and the apoplast (cell walls) (Buckley & Sack, 2019; Rockwell *et al.*, 2022; Scoffoni *et al.*, 2023; Jain *et al.*, 2024). Based on the calculations of Buckley & Sack (2019),

the lowest reported intercellular RH of c. 0.7 in our study corresponds to a water potential of c. -50 MPa. Given that a 1% decline in intercellular RH reduces water potential by c. 1.4 MPa (Buckley & Sack, 2019), very low-water potential values could occur in a range of RH levels. It is difficult to know at which location inside the leaf tissue such extraordinarily negative water potentials would occur, considering that water vapour potentially travels from the evaporation sites deeper into the mesophyll air spaces through the substomatal cavity to reach the stomatal pore (Wong et al., 2022). Nevertheless, a bulk leaf water potential value this negative would be lethal for most plants (Bartlett et al., 2012, 2016; Klein, 2014). However, we did not observe any sign of wilting in any of the tested species throughout the measurements. Therefore, we must assume that: (1) the extreme negative water potential is restricted to the apoplast and (2) the leaf water potential in living cells and bulk leaf water potential are much less negative than that in the apoplast.

These assumptions imply that the plasma membranes would be far less permeable to water at high VPD to sustain such a substantial difference in the water potential between the apoplast and inside the living cells (Wong *et al.*, 2022; Scoffoni *et al.*, 2023; Cernusak *et al.*, 2024). According to the estimation of Wong *et al.* (2022), this permeability is indeed expected to be lower than values reported in previous studies (Morillon & Chrispeels, 2001; Martre *et al.*, 2002). However, as *E* increased with increasing VPD (Fig. 1b), other water flow pathways could be dominant to sustain the water supply to the evaporation surfaces at high VPD. The symplastic pathway through plasmodesmata could be responsible for sustaining water movement between cells, although this is still under debate (Fricke, 2000; Barbour & Farquhar, 2004).

Although the plasma membrane potentially has sharply reduced conductance to water at high VPD, it apparently can largely maintain its conductance for CO2 transport. In support of this perspective, at the highest VPD, c_c/c_a showed a small reduction when saturation was assumed (Fig. 5c) and a moderate reduction when unsaturation was accounted for (Fig. 5f), compared with c_i/c_a (Fig. 5a,d), suggesting a high plasma membrane conductivity for CO_2 at high VPD. The membrane conductance to CO_2 can be facilitated by some aquaporins (Uehlein et al., 2008; Zhao et al., 2017; Chen et al., 2023). Moreover, it has been hypothesized that the aquaporin-mediated regulation of diffusion can be different for CO₂ and H₂O (Otto et al., 2010; Kaldenhoff, 2012). It has been proposed that aquaporin subunits with different affinities to CO₂ and H₂O compete in the formation of aquaporin complexes, which reduces water permeability when CO2 conductivity through aquaporins is high (Flexas et al., 2012; Hommel et al., 2014). Certainly, more studies are required to reconcile the mismatch between unsaturation of leaf air spaces and leaf water potential. This has strong implications for improving our understanding of leaf water transport mechanisms and could potentially change our fundamental understanding of leaf CO₂ and H₂O exchange.

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Competing interests

None declared.

Author contributions

HD, LAC, MML and RTWS conceived the study. MML supervised the project. HD and RTWS set up the instrumentation and carried out the experimental work. HD and LAC processed the experimental data and performed the analysis. HD wrote the manuscript, whilst LAC, MS, AG, RTWS and MML contributed critically to the manuscript.

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Data availability

All data used in this study are available in the Supporting Information Dataset S1.

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Supporting Information

Additional Supporting Information may be found online in the Supporting Information section at the end of the article.

Dataset S1 The dataset analysed in this study.

Fig. S1 Variations in air relative humidity (RH) with increasing air vapour pressure deficit (VPD).

Fig. S2 Online stable carbon ($\Delta^{13}C_{obs}$) and oxygen ($\Delta^{18}O_{obs}$) discriminations of CO₂ during assimilation in response to increasing air vapour pressure deficit (VPD).

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Fig. S3 Response of leaf intercellular relative humidity (e_i/e_s) to increasing air vapour pressure deficit (VPD).

Fig. S4 Responses of intrinsic water use efficiency (iWUE) to increasing air vapour pressure deficit (VPD).

Fig. S5 Impacts of leaf intercellular relative humidity (e_i/e_s) and leaf conductance (g_t) on the potential transpiration rate (E) with increasing air vapour pressure deficit (VPD).

Methods S1 Supplementary equations for calculations of intercellular relative humidity from $\delta^{18}O.$

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