1	Framing.	Australian	Pleistocene	Coastal	Occupation	and A	Archaeol	logy
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17 Abstract

There are few archaeological sites that contain records for Pleistocene coastal occupation in Australia, as is the case globally. Two major viewpoints seek to explain why so few sites exist. The first is that the Pleistocene coast was a relatively marginal environment where fluctuating sea levels actively inhibited coastal resource productivity until the mid-to-late Holocene. The second position suggests that the Pleistocene coast (and its resources) was variably productive, potentially hosting extensive populations, but that the archaeological evidence for this occupation has been submerged by sea level rise. To help reconcile these perspectives in Australia, this paper provides a review, discussion, and assessment of the evidence for Australian Pleistocene coastal productivity and occupation. In doing so, we find no reason to categorically assume that coastal landscapes were ever unproductive or unoccupied. We demonstrate that the majority of Pleistocene coastal archaeology will be drowned where dense marine faunal assemblages should only be expected close to palaeo-shorelines. Mixed terrestrial and marine assemblages are likely to occur at sites located >2km from Pleistocene shorelines. Ultimately, the discussions and arguments put forward in this paper provide a basic framework, and a different set of environmental expectations, within which to assess the results of independent coastal research.

32 Keywords: Pleistocene, coastal archaeology, Australia, sea levels

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34 1 Introduction

35 Relatively little is known about occupation of Pleistocene coastal landscapes in Australia. This is partly 36 because, compared to terrestrial Pleistocene records, there are few archaeological sites that contain records 37 for Pleistocene coastal occupation. There are two major viewpoints that explain why so few sites contain records for Pleistocene coastal occupation around the world. The first is that the Pleistocene coast was a 38 39 relatively marginal environment where fluctuating sea levels actively inhibited coastal resource productivity 40 (e.g. Beaton, 1985; Cohen, 1977; Mulvaney and Kamminga, 1999; O'Connell et al., 2012; Osborn, 1977; 41 Yesner, 1987; see Bicho and Hawes, 2008:2166-2167 and Erlandson, 2001 for summaries). According to 42 this perspective it was only late in human history, under relative mid-to-late Holocene sea-level stabilisation, that humankind began to intensively exploit coastal resources, producing a 'coastal adaptation'. The second 43 44 position argues that the Pleistocene coast (and its resources), was variably productive and critically 45 important in facilitating both the evolution and dispersal of humankind (Bailey et al., 2007, 2015; Erlandson, 46 2001; Erlandson and Braje, 2015; Erlandson et al., 2007, 2015; Kyriacou et al., 2014; Marean, 2014; 47 Parkington, 2010; Ward et al., 2015; Will et al. 2016). According to this position, rising post-glacial seas 48 drowned most Pleistocene evidence for coastal occupation creating a preservation bias towards mid-to-late 49 Holocene coastal archaeology (e.g. Bailey and Milner, 2002; Bailey and Flemming, 2008; Bicho and Hawes, 50 2008; Bicho et al., 2011; Erlandson, 2001; Erlandson and Fitzpatrick, 2006). Nevertheless, regardless of which argument is privileged, the problem still remains: we know very little about how people occupiedand used now-drowned Pleistocene coasts and coastal landscapes.

53 The coastal archaeological record in Australia is, like many areas of the world, dominated by a mid-to-late 54 Holocene coastal archaeology (e.g. Rowland et al., 2015; Ulm, 2011), although there was some recognition 55 for the early use of coasts (e.g. Bowdler, 1977, 1990; Dortch, 1997; Morse, 1993a; White and O'Connell, 56 1982). The lack of Pleistocene evidence has been interpreted in light of the first viewpoint mentioned above: 57 coasts were relatively unproductive due to fluctuating sea levels and only become economically important 58 for Aboriginal hunter-gatherers from the mid-Holocene when stabilised sea levels resulted in widespread 59 coastal productivity (e.g. Beaton, 1995; Hiscock, 2015; Nunn, 2020; Mulvaney and Kamminga, 1999; 60 O'Connell and Allen, 2012; Pope and Terrell, 2007). While some evidence, particularly from the coastal 61 northwest and New Ireland has challenged this view (e.g. Allen et al. 1989; Morse, 1999; Veth et al., 2007, 62 2017c; Ward et al., 2015), Pleistocene coastal archaeology is rare throughout the continent. This ultimately 63 leaves two perspectives which seek to explain the relative lack of Australian Pleistocene coastal archaeology and so generates two important related questions. First, were Australian Pleistocene coasts productive 64 65 landscapes? Second, did people widely occupy the varied Pleistocene coastal landscapes throughout 66 Australia?

To help answer these questions, this paper reviews, discusses and assesses the evidence for Australian 67 Pleistocene coastal productivity and occupation, including evidence from the wider Sahul region. However, 68 69 these are difficult questions to answer because almost all relevant landscapes are now underwater. The available archaeological record on islands, archipelagos and areas adjacent to steep continental shelves 70 71 represent the most tangible links to drowned coasts and so are the primary datasets that have been used to 72 reconstruct the interaction of people with Australian Pleistocene coasts. In the first part of this paper, we 73 review this archaeological literature to outline the current state of knowledge on Australian Pleistocene 74 coastal archaeology. This review will outline the current standing of the Australian coastal literature and 75 form a baseline to begin answering the two questions posed above. Drawing on international and 76 palaeoenvironmental literature, in the second part of the paper we assess Australian Pleistocene coastal 77 productivity and the likelihood of coastal occupation. This discussion provides a framework within which

78 research questions and agendas can be better situated to study Pleistocene coastal occupation in Australia.

79 Based on this framework, we identify major directions for future research on Pleistocene coastal landscapes.

80

81 2 Defining the Coast and its Resources

82 Since coastal landscapes support a variety of environments, it is important to set out some general 83 definitions. This is especially important because the terms 'coast(al)', 'marine' and 'maritime' are often used 84 interchangeably with only passing reference to their specific meaning (Bailey et al., 2015:44; Erlandson and Fitzpatrick, 2006). In recognising these problems, recent work has sought to provide clearer definitions. 85 86 Erlandson and Fitzpatrick (2006:8-9) differentiate between 'coastal' and 'maritime' adaptations. They define 87 'coastal adaptations' as 'any subsistence lifestyle based along the margins of a large body of water that includes 88 the regular use of foods from aquatic habitats' while 'maritime adaptations' are defined as 'those cases where 89 humans regularly used boats for travel and subsistence purposes, where voyaging away from the immediate 90 coastline was possible, and where the majority of nutrition (calories or protein) was derived from marine 91 resources' (Erlandson and Fitzpatrick, 2006:8-9). Marean (2014:20) suggests a similar set of definitions (see 92 also Jeradino, 2016b; Will et al., 2016, 2019). However, these definitions tend to focus on ecotones associated with, or seaward of, the shoreline. The 'coast' also exerts a significant influence on ecotones 93 94 located immediately landward of the shoreline. As Bailey et al. (2015) have pointed out, the term 'coastal' 95 can refer to a large region of variable extent, possibly extending many kilometres inland. This landscape is 96 often referred to as the 'coastal plain'. It contains ecotones under the influence of both terrestrial and 97 marine processes, often encompassing flora and fauna unique to this zone and economically attractive to 98 hunter-gatherers (see below). Yet, simply lumping the coastal plain in with the 'marine' (or even 'maritime') 99 terms risks creating a generalisation so broad that it lacks any conceptual usefulness (Bailey et al., 2015; 100 Hallam, 1977).

In light of these considerations, we suggest slightly revised definitions. First, following Erlandson and
Fitzpatrick (2006:8-9) and Marean (2014:20), 'maritime' here refers to the zone, and all resources, that lie
to seaward of the intertidal zone beyond pedestrian foraging accessibility that require watercraft technology
for access (see also Fa, 2008:2203-2204). 'Maritime adaptations' are defined by economic accessibility to

105 this zone but importantly can also be defined by archaeological technology (e.g. Balme, 2013) or cultural 106 (e.g. McDonald, 2015) indicators for access to the maritime zone where the majority of nutrition does not 107 necessarily have to derive from maritime sources (c.f. Erlandson and Fitzpatrick, 2006). 'Intertidal' refers 108 to the area of coastline periodically exposed and inundated by tides but, importantly, are accessible by 109 pedestrian foraging and do not require watercraft technology. 'Marine' is used more generally and refers to 110 the zone permanently or tidally submerged by ocean waters incorporating both the maritime and intertidal 111 zones. The 'coastline' or 'shoreline' is used to refer to the narrow zone where land meets the sea (but see 112 Larcombe et al. 2018). The 'coastal plain' refers to all those zones occurring landward of the shoreline where ecosystems are significantly influenced by proximity to the ocean or other variables unique to this 113 114 zone (e.g. slope, exposure to on-shore winds etc.). Finally, 'coast' is used more generally here, referring to all those zones, which uniquely occur due to oceanic proximity, and that are accessible by pedestrian 115 116 foraging. The 'coast' then subsumes the coastal plain, the shoreline and the intertidal zone. These definitions, and their wider discussion in the literature, ultimately show that the maritime, marine, coastal 117 and intertidal zones are spatially and temporarily dynamic zones which resist easy definition. 118

119 The coast can be an extraordinarily productive and resource-rich environment. Coastal gross primary 120 productivity is around 2000 kcal/m²/yr, which is twice as productive as the open ocean, while the primary 121 productivity of estuarine and intertidal zones can reach 10 times this level (Fa, 2008:2195, and references 122 therein; Woodroffe et al., 1988, 1989). However, the coast incorporates more than just marine resources. 123 It is a 'super-ecotone' where a mosaic of terrestrial and marine habitats co-exist over small distances, often 124 including freshwater environs in the form of rivers, creeks, springs and seeps (Bailey et al., 2007, 2008; 125 Bailey and King, 2011:15; Barker, 1999; Bicho et al., 2011:xviii; Erlandson and Braje, 2015:34; Erlandson et al. 2015; Hallam 1987; Perlman, 1980:281). Although coasts tend to be relatively productive 126 127 environments, they are not uniformly productive across space or time (Erlandson, 2001:331-332). Instead, a variety of processes operating on different scales dictate the structure of coastal landscapes. Larcombe et 128 al. (2018) provide a detailed review of the fundamental physical processes that define coastal morphology 129 130 and change across space and time. For example, slope, tidal amplitude, wave action, bathymetry, sedimentation (erosional or depositional), oceanographic variables (e.g. sea level fluctuations, dominant 131 132 currents, primary productivity, upwelling, salinity and sea surface temperature), climate and tectonics can

all significantly structure coasts and their relative productivity (see also Bailey and King, 2011; Bailey et al.,
2007; Bird, 2008; Chappell and Thom, 1977:278-279; Davidson-Arnott, 2010; Fa, 2008; Fa and Sheader,
2000; Lambeck and Chappell, 2001; Lambeck and Nakada, 1990; Lambeck et al., 2002, 2014; Lewis et al.,
2011; Jennings, 1971; Murray-Wallace and Woodroffe, 2014; Perlman, 1980; Pye and Allen, 2000;
Semeniuk, 1995; Siddal et al., 2003; Webster et al. 2018). Indeed, these processes have left complex
signatures on now-drowned coasts with varying consequences for the preservation of archaeological
deposits (e.g. Brooke et al., 2017).

140 Sea level fluctuation, being the most prominent temporal coastal process throughout the Pleistocene, 141 deserves some expanded discussion. Indeed, between 18,000 and 8,000 years ago sea levels around the Australian margin rose 130m (Ishiwa et al., 2016; Lewis et al., 2013). There can be no doubt that long-term 142 143 Pleistocene sea level fluctuations resulted in localised coastal alteration or re-organisation including a 144 constant dynamic flux in productivity, resource structure and the organisation of occupation (Bailey and 145 King, 2011; Bird, 2008; Chappell and Thom, 1977; Davidson-Arnott, 2010; Erlandson and Fitzpatrick, 2006; Hepp et al., 2019; Hinestrosa et al., 2016, 2019; Jennings, 1971; Larcombe et al. 2018; Reeder-Myers 146 147 et al., 2015; Williams et al., 2018; Woodroffe, 1990; Woodroffe et al., 1988:96-97). In Australia, sea level 148 regime (i.e. regression or transgression) has been argued to play a major role in determining coastal productivity and configuration (see Allen et al., 2020; Beaton, 1985; Bowdler, 2010; Chappell, 1993, 2000; 149 Chappell and Thom, 1977; Grindod et al., 1999; Hall, 1999; O'Connell et al., 2010). Regressive phases have 150 151 been associated with river entrenchment, swamp contraction, modest mangrove development, saline flats 152 and lower levels of productivity where sediments tend to accumulate in alluvial valleys instead of the shore 153 and/or bypass the estuarine system altogether due to river entrenchment. Transgressive phases have been 154 associated with more productive estuaries, lagoons and coral reefs where sediment often becomes trapped 155 within the coastal system, contributing towards the expansion of swamps and estuarine environments (e.g. 156 Grindrod et al., 1999; Johnson et al., 1982). Of course, the exact influence of any regressive or transgressive regime will be dependent on local coastal structure and the pace and magnitude of sea level change 157 158 (Hinestrosa et al., 2016, 2019; Larcombe et al., 2018; Ward et al., 2015; Webster et al., 2018). One well-159 studied example of coastal dynamism under a transgressive regime can be sourced from work on Australian 160 mangroves in the Alligator Rivers region of the Northern Territory, which exhibit extraordinary diversity

in their responses to early-to-mid-Holocene sea level rise (see Chappell and Thom, 1977; Clark and Guppy,
1988; Grindod et al., 1999; Jennings, 1971; Proske et al., 2014; Semeniuk, 1983, 1994, 1995; Thom et al.,
1975; Ward et al., 2015; Wolanski and Chappell, 1996; Woodroffe, 1990; Woodroffe et al., 1985, 1988,
1989, 1993). Ultimately, although sea level fluctuation is a global phenomenon, its impact on coastal
environments and resident populations varies locally because all coastal settings interact with sea level
fluctuation uniquely.

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168 3 Pleistocene Coastal Occupation in Australia: A Review of Evidence, Significance and Issues

169 3.1 Earliest Evidence and Coastal Colonisation

170 It is widely agreed that Sahul was settled by 'behaviourally modern' people sometime before 50,000 years ago (Balme, 2013; Balme et al., 2009; Bradshaw et al., 2019, 2021; Clarkson et al., 2015, 2017; Hiscock, 171 172 2013, 2015; Miller et al., 2016a, 2016b; O'Connell et al., 2018; Smith, 2013; Veth et al. 2017c). Sahul sits at 173 the end of the Southern Dispersal Route and its settlement required maritime (specifically, watercraft and 174 cordage) technology to make a water crossing of up-to 120km indicating that, not only was the settlement 175 of Sahul a coastal one, but that the settling population also possessed a coastal and maritime adaptation 176 (Balme, 2013; Balme et al., 2009; Bird et al., 2018, 2019; Birdsell, 1977; Bowdler, 1977; Chappell, 2000; 177 Hiscock, 2013; Jones, 1979; Kealy et al. 2016, 2017; Kuijjer et al. 2022; O'Connell and Allen, 2012, 2015; 178 O'Connell et al., 2010; O'Connor and Chappell, 2003; O'Connor and Veth, 2000; Szabo and Amesbury, 179 2011). The evidence from Sahul and its nearby Pleistocene islands is, outside of Africa, among the earliest 180 coastal archaeological evidence associated with Homo sapiens and provides some of the only conclusive 181 evidence for coastal and maritime adaptations on the Southern Dispersal Route (Barker 2013; O'Connor 182 et al. 2011; Leavesley and Allen 1998; Veth et al., 2017c). Indeed, there is currently very little archaeological 183 evidence for pre-glacial coastal occupation between eastern Africa and South East Asia, leaving a 184 >10,000km long gap along the Southern Dispersal Route with supporting evidence only from the beginning 185 (e.g. Marean et al. 2011) and the end (Bailey et al., 2015; Bulbeck, 2007; Erlandson and Braje, 2015; but see 186 Walter et al., 2000; cf. Bailey and Flemming, 2008:2156; Bailey et al., 2007:146-147, 2015:52). Whichever 187 routes were taken, the evidence shows that, at least by the time people reached Wallacea, they were

exploiting coastal resources (e.g. Leavesley et al., 2002; O'Connor, 2007; O'Connor et al., 2011). Recent
modelling shows that the now submerged coastal margins were likely to have been key corridors of
movement for the earliest populations (Crabtree et al., 2021).

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192 3.2 A Coastal Time Lag in Australia?

193 Even though maritime adapted people began the settlement of Sahul, the Australian evidence for 194 Pleistocene coastal occupation and resource use is rare until the mid-to-late Holocene (Bowdler, 1995; 195 Chappell, 2000; O'Connell et al., 2010; O'Connor and Veth, 2000; Richards 2012; Rowland et al. 2015; 196 Ulm, 2011, 2013). As Barker (1999:119) noted, 90% of the dated coastal sites in Australia only retain 197 evidence for mid-to-late Holocene coastal occupation (e.g. Ulm 2011). This archaeological phenomenon 198 was initially used to argue that Pleistocene to early Holocene sea level fluctuations inhibited coastal 199 productivity until mid-to-late Holocene sea level stabilisation allowed productive conditions to develop 200 facilitating coastal occupation (Beaton, 1985; Callaghan, 1980; Hughes and Lampert, 1982; Lampert and 201 Hughes, 1974; Mulvaney and Kamminga, 1999; Rowland, 1983, 1999; Walters, 1989). Perhaps the most 202 well-known example is Beaton's (1985) time-lag hypothesis based on research at Princess Charlotte Bay in north Queensland. The absence of shell middens pre-dating approximately 4,700 BP led Beaton to conclude 203 204 that coastal occupation did not occur until 1500 years after the marine transgression had stabilised. Beaton 205 attributed this to the post-glacial marine transgression which, he argued, prevented productive coastal 206 ecosystems from forming and that, even following sea level stabilisation (6000 years ago), coastal 207 environments 'lagged behind' as they slowly regained productivity. Once productivity levels were sufficient, 208 coasts were occupied by Aboriginal people resulting in the mid-to-late Holocene coastal archaeological 209 record (Beaton, 1985).

This argument was quickly rebutted by other Australian researchers in both Western Australia and Queensland (Barker 1999, 2004). In northwest Australia both Veth (1993; Veth et al., 2007) and Morse (1988, 1993a, 1993b; see also Przywolnik, 2002) found evidence for Pleistocene to early Holocene coastal economies. Morse (1993a; see also O'Connor, 1999; Veth, 1999) suggested that, despite changes in sea level, coastal resources were always part of past Aboriginal economy and, as sea levels fluctuated, people

followed the sea. In the coastal Kimberley, evidence from Koolan Shelter 2 also demonstrated that marine 215 resource use was well established during the terminal Pleistocene (O'Connor, 1999). In Tasmania, Rocky 216 217 Cape South and Cave Bay Cave both contain relatively dense shell middens that include fish remains dating 218 to around 7000 – 8000 BP as the sea rose to its current position (Bowdler, 2010; Jones, 1968). Furthermore, off the coast of Queensland on the Whitsunday Islands, Barker (1991, 1999, 2004) demonstrated marine 219 220 resource use from archaeological deposits at Nara Inlet 1 and Border Island 1 in association with sea level 221 rise. Barker suggests that this provides good evidence for both the resilience of Aboriginal coastal 222 occupation and productive marine environments in the face of sea level transgression (see also Rowland et al. 2015). Although, as noted by Rowland et al. (2015:158) and Rowland et al. (2021), McNiven et al. (2014) 223 224 have recently recast the time-lag hypothesis to explain delayed settlement patterns on islands in Shoalwater 225 Bay in Queensland.

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227 3.3 Models for Australian Pleistocene Coastal Occupation and Productivity

228 While some further evidence for Australian Pleistocene marine resource use has since been found (e.g. Przywolnik, 2002; Richards 2012; Veth et al., 2007), on a continental scale, evidence for Pleistocene coastal 229 occupation remains rare and ephemeral. Bowdler (2010) has even suggested that much of the eastern 230 231 Australian coast was relatively unproductive and unattractive during the Pleistocene resulting in little 232 occupation (see also Mulvaney and Kamminga, 1999). As a result, many scholars have argued that the early 233 settlement of Australia largely occurred in the interior where coasts were subsequently populated from 234 savannah landscapes (e.g. Chappell, 2000; Hiscock, 2008, 2015; Hiscock and Wallis, 2005; O'Connor and Chappell, 2003; O'Connor and Veth, 2000; Smith, 2013; see also Hallam, 1987; Horton, 1981). In this 235 236 literature, marine resources are generally perceived as *ad hoc* additions to, or part of a more generalised 237 mixed, economy and do not become important until the Holocene (e.g. O'Connor and Veth, 2000; but see 238 Morse, 1993a, 1999). In recognising the emergent terminal Pleistocene evidence for coastal occupation, Beaton (1995:798-802) proposed a dichotomous model that distinguished procumbent coastlines (gentle 239 240 slope, low relief and low wave energy environments) from precipitous coastlines (high relief and slope with 241 high wave energy). During the last marine transgression, the latter should be more productive because of their relative stability. Beaton (1995) rejected the sparse and terrestrially mixed Pleistocene marine resource
archaeological assemblages as evidence for a 'coastal economy', reiterating that coastal economies
(archaeological assemblages dominated by marine fauna) are only a mid-to-late Holocene phenomenon
under 'highly productive' stabilised coastal environments (see also Smith, 2013; Hiscock, 2015).

Beaton's (1995) dichotomous model also formed part of a colonisation model for Australia by O'Connell 246 247 and Allen (2012, 2015; Allen and O'Connell, 2020; O'Connell et al., 2010; but see Veth et al. 2007). Upon 248 reaching Sahul, O'Connell and Allen (2012:7) suggest that coasts were used in association with movement 249 into the interior but that precipitous and procumbent coastal patches were ranked differently (O'Connell 250 and Allen, 2012:8). They argue that shellfish on precipitous shorelines will redistribute quickly in response 251 to sea level change because of their steep slopes and rocky substrates but, because precipitous intertidal 252 zones are narrow, high ranked prey species were quickly depleted by predation (see also Codding et al. 253 2014). In contrast, O'Connell and Allen (2012:8) argue that shellfish on procumbent shorelines, although 254 locally abundant and resistant to over-predation, were devastated by sea level change, with overall 255 productivity only recovering well after sea level stabilisation (sensu Beaton, 1985). They argue that although 256 precipitous shorelines are highly ranked patches, their potential to be rapidly depleted meant that more 257 permanent occupation of areas adjacent to precipitous coastlines only occurred after Holocene sea level stabilisation (sensu Beaton, 1985, 1995). As most Pleistocene and early Holocene archaeological coastal sites 258 259 have been recorded from precipitous coastlines and contain sparse marine fauna (e.g. Morse 1993a; 260 Przywolnik, 2002), O'Connell and Allen (2015:76) suggest their predictions are supported. As such, they 261 argue that Australia was largely settled via the interior where high-ranked terrestrial patches were targeted.

262 However, the coastal component of O'Connell and Allen's (2012) model has been criticised due to its over-263 emphasis of the vulnerability of coastal productivity due to sea level fluctuations, especially on 'procumbent 264 coasts' (Ditchfield et al., 2018; Erlandson, 2012; Manne and Veth, 2015; Veth et al., 2014, 2017b, 2017c; 265 Ward et al. 2014, 2015). These critiques have proposed that Pleistocene marine ecosystems were productive 266 and able to respond quickly to sea level change, being in dynamic equilibrium with the coast if sediments 267 were available. It is suggested that this was especially the case for 'procumbent coasts', since broad, low-268 relief, coastal plains likely provided significant opportunity for sediment accumulation resulting in 269 productive coastal environments with abundant resources for coastal foragers. Veth et al. (2014) cite 270 evidence for continued marine resource use from transgressing, relatively broad and low-relief ('procumbent') coasts since 14,000 cal. BP in northwest Australia. Richards (2012) work at Cape Duquesne 271 272 in southwest Victoria provides further evidence for early Holocene marine resource use with a variety of 273 coastal shell middens returning dates from 11600 - 8600 cal. BP, while Nunn (2020) has recently shared 274 several Aboriginal histories about people interacting with the post-glacial marine transgression. 275 Assumptions about the vulnerability of large shellfish to predation are also open to challenge based on 276 results showing long-term sustainability of shellfishing in areas with high productivity but likely low 277 populations (e.g. Ulm et al., 2019). It also is worth noting that Pleistocene sea level fluctuation was not always constant and that sea levels may have stabilised under productive conditions for short periods. For 278 279 example, O'Leary et al. (2020) recently identified a relict drowned MIS3 shoreline from a previous stabilised 280 period with potentially productive components such as estuaries and lagoons (see also Brooke et al., 2017).

281 In response to some of this literature, O'Connell and Allen (2015:76) have argued that sedimentary 282 substrates would often require centuries to re-form as a result of sea level fluctuation and that productive coasts, especially the 'procumbent' type, were rare until mid-Holocene sea level stabilisation. They reaffirm 283 284 that the correlation of early 'near-coastal' sites adjacent to steep shelves fits with their predictions whereby 285 only precipitous coasts will be productive and therefore attractive to early populations (O'Connell and 286 Allen, 2015:76). Williams et al. (2018:151) have also suggested that, since the precipitous northwest coastal 287 sites are constrained by desert or arid environments, there was little other alternative beyond marine 288 resources for the past occupants of these coastal locations.

289 It is worth noting here, that procumbent coasts supported many more resources than simply shellfish. 290 Hallam (1987) originally made this point in response to Bowdler's (1977) coastal colonisation model, 291 arguing that coasts include significant hinterland habitats such as swamps, lakes, floodplains and savannah, 292 and that these likely contributed a significant proportion of coastal diet. Importantly, this also includes 293 terrestrial plant resources which, due to their limited preservation in the archaeological record, often receive 294 less attention in reconstructions of past coastal diets (Roberts et al., 2020). Veth et al. (2007, 2014, 2017b; 295 Ditchfield et al., 2018) have repeatedly stressed the existence of such a broad-based economy for the 296 northwest coastal plain for over 42,000 years despite sea level fluctuation. Indeed, during a review of important Australian palaeoclimatic records, De Deckker et al. (2020:24) have also suggested that thePleistocene coast "would have nurtured more human activity during the settlement of this landmass".

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300 3.4 Southeast Asian Evidence from Sahul

No review of Pleistocene coasts for Sahul would be complete without the evidence from southeast Asia. 301 302 Freshwater resources appear to be the earliest exploited aquatic fauna, with remains of freshwater molluscs 303 from the sites of Kao Pah Nam in Thailand (700,000 BP) and Trinil in Java (500,000 BP to 400,000 BP) 304 thought to have been procured by Homo erectus (Choi and Driwantoro, 2007; Joordens et al., 2014; Ono, 305 2016). In terms of Anatomically Modern Humans, Szabo and Amesbury (2011) noted that early sites in this 306 region are also characterised by freshwater rather than marine mollusc exploitation. For example, at Niah 307 Cave in Borneo, modest quantities of freshwater molluscs have been recovered dating back to 50,000 BP, 308 along with the remains of freshwater turtles (Piper and Rabett 2014; Szabo and Amesbury 2011). Later in 309 the sequence, the inclusion of greater quantities of estuarine taxa attest to shifting coastlines in the early 310 Holocene (Szabo and Amesbury 2011).

311 Early evidence for the exploitation of marine resources have been recovered from Laili Cave (44,000 BP), Asitau Kuru (formerly known as Jerimalai, 42,000 BP), Gua Makpan (40,000 - 38,000 BP), Buang Merabak 312 313 (41,000 BP) and Matenkupkum (41,000 BP). These assemblages display either use of near-shore marine 314 resources and/or the exploitation of pelagic fish (e.g. Scombrids exploited between 42,000 - 38,000 BP at 315 Asitau Kuru) since occupation (Hawkins et al., 2017; Kealy et al. 2020; Leavesley and Allen, 1998:75; O'Connell et al., 2010: 60; O'Connor and Chappell, 2003:17; O'Connor et al., 2011, 2017a; but see 316 317 Anderson, 2013a, 2013b; Bailey, 2013; Erlandson, 2013; cf. O'Connor and Ono, 2013 for debate about 318 Asitau Kuru).

Dating to 37,000 BP, the earliest evidence for the manufacture of shell beads in Southeast Asia has been located from Asitau Kuru, in the form of an *Olivia* bead (Langley et al. 2016). The preference for, and use of, *Olivia* beads is attested by their presence not only throughout the Asitau Kuru sequence, but also from the nearby sites of Lene Hara and Matja Kuru 1 and 2. Some of the earliest evidence for complex fishing technology worldwide, is found from the sites of Matenbek and Asitau Kuru, dating to 22,000 BP (Allen et

al., 1988; O'Connor and Ono, 2013; Smith and Allen, 1999; c.f. Langley et al. 2021; but see Yellen et al., 324 1995). At Gua Makpan, the recovery of 239 specimens relating to fishing technology and dating to the last 325 326 15,000 years, demonstrates the importance and wide-ranging use of marine shell for these activities (Langley 327 et al. 2021). Ornamental artefacts were also manufactured at Gua Makpan (Kealy et al. 2020) and Asitau 328 Kuru (Langley et al. 2016), from Nautilus shell. (see also Langley and O'Connor, 2017 and Langley et al., 329 2019 for a review). Langley et al. (2016) suggest that Asitau Kuru's archaeological record, with its production 330 of ornamental artefacts and manufactured fishing technology, indicates a coastal adaptation in which the 331 coastal landscape was intertwined with the social realm. The recovery of fish-hook technology in association 332 with a terminal Pleistocene burial at Tron Bon Lei (Alor Island), further demonstrates the social and 333 cosmological connection that the coast and its marine resources had for the Pleistocene inhabitants of Southeast Asia (O'Connor et al., 2017b). 334

335 Like many sites worldwide with Pleistocene coastal signatures (e.g. Erlandson, 2001), Szabo and Amesbury (2011) note that many of these sites are located next to steep bathymetry where most sites are likely located 336 337 within 1km of LGM shorelines. For example, evidence of pre-Holocene exploitation of precipitous coasts 338 can be found in northern New Guinea, along the steep Vanimo coast in West Sepik province (Gorecki et 339 al., 1991; O'Connor et al., 2011). Watinglo and Lachitu Rockshelters contain evidence of terminal Pleistocene/Holocene coastal use, although early dates of 30,444 - 29,380 and 29,065 - 28,000 cal. BP 340 341 obtained from marine shellfish in Lachitu Rockshelter imply this site may have had a longer occupation 342 sequence (O'Connor et al., 2011:9). Both sites contain dense zones of marine shellfish, with the Lachitu 343 assemblage representing a diverse suite of habitats, including rocky, sandy, reef and mangrove (Gorecki et 344 al., 1991; O'Connor et al., 2011; also see Summerhayes et al., 2017 for further review). However, given the 345 limited nature of overall Pleistocene coastal exploitation, Szabo and Amesbury (2011: 12) question whether coastal ecosystems, especially estuarine ones, were ever stable in the region during Pleistocene sea level 346 347 fluctuation. Following Terrell (2004), Allen and O'Connell (2020) have also suggested that reefs, lagoons, 348 swamps and floodplains along the northern New Guinea coast were replaced by unproductive rocky coasts 349 with entrenched rivers during the glacial, an argument that they suggest can be extended to Sahul's steep rocky coasts. They suggest this environmental shift may have acted to restrict mobility and isolate Sahul 350 351 from Wallacea during the LGM. Following the LGM and from the terminal Pleistocene onward, there is a

rapid rise in marine resource use represented in sites throughout the region (transgression and stabilisation;
see Ono et al., 2020; Szabo and Amesbury, 2011), which includes evidence for pelagic fishing and complex
maritime technology especially on islands depauperate in terrestrial fauna (e.g. Carro et al., 2016; Kealy et
al., 2020; O'Connor et al., 2018).

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357 3.5 Remarks on Pleistocene Coasts in Australia

358 Clearly, many issues remain for Pleistocene coastal archaeology in Australia but there are perhaps two especially pertinent questions. First: why is there still so little evidence for Pleistocene coastal occupation 359 360 in Australia? The notion that Pleistocene coasts were largely unproductive and only supported ephemeral 361 occupation on precipitous shorelines provides one possible answer. Sea level rise drowning most of the evidence for Pleistocene coastal occupation provides another. This brings us to the second pertinent 362 question: independent of occupation, were Pleistocene coasts productive? This review has shown that the 363 364 nature of Pleistocene coastal productivity is contested. Building from this review and drawing on 365 international and national palaeo-environmental literature, we now discuss, analyse and assess the current 366 evidence for Australian Pleistocene coastal productivity and occupation.

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4 Regional Pleistocene – Early Holocene Coastal Occupation and Productivity in Australia: A Discussion

This discussion has two major aims. First, to assess how much of the coastal record was drowned and, from
that discussion, assess the representativeness of the record that remains. Second, to assess whether
Australian Pleistocene coasts were productive based on current coastal palaeoenvironmental and
archaeological research.

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377 4.1 How Representative is the Coastal Record?

Many scholars have argued that most pre-Holocene coastal zones were drowned, eroded or buried by rising 378 seas during the terminal Pleistocene, thereby making earlier coastal records inaccessible and creating a false 379 380 impression of a global mid-to-late Holocene 'coastal efflorescence' (Bailey and Flemming, 2008; Bailey and 381 Milner, 2002:4; Bailey et al., 2015; Bicho and Hawes, 2008; D'Alpoim Guedes et al., 2016; Erlandson, 382 2001:300; Erlandson and Fitzpatrick, 2006; Parkington 1980; Rick et al., 2005:176). Indeed, as Bailey et al. 383 (2007:130-131, 138, 2015:43) point out, for most of human existence sea levels generally oscillated between 384 -40m to -60m below present levels, meaning that, if people occupied these coastal zones, the majority of 385 pre-Holocene coastal archaeology will have since been submerged. Therefore, at the global scale, the largely 386 terrestrial archaeological record may not be fully representative of human evolution, dispersal, and 387 environmental interaction and may provide a biased, perhaps misleading and incomplete picture (Bailey and Flemming, 2008:2157; Bailey and Milner, 2002:4-5; Bailey and King, 2011:15; Bailey et al., 2007:131; 388 389 Erlandson, 2001; Erlandson and Braje, 2015; Erlandson and Fitzpatrick, 2006:6; Perlman, 1980:296). This 390 may also be true for the Australian record.

391 This argument assumes that sea level rise was significant enough to submerge most archaeological deposits 392 that would otherwise provide evidence for widespread Pleistocene coastal occupation. This assumption 393 should be tested on a case-by-case basis where the question could be framed as: under conditions of 394 significant Pleistocene coastal occupation, should there be conclusive evidence for coastal occupation 395 preserved in terrestrial contexts today? If the answer is 'yes', then the relative lack of evidence for Pleistocene coastal adaptations may be a genuine behavioural phenomenon (or perhaps the product of 396 397 limited sampling). If the answer is 'no', then we must confront the fact that the archaeological record under-398 represents coastal occupation and consequently consider what a Pleistocene coastal adaptation might look 399 like. One way to begin addressing the question for Australia is to consider the average distance within which 400 archaeological evidence for coastal occupation is likely to occur from palaeoshorelines. There is good 401 evidence to consider that distance from palaeoshorelines may play a significant role in the archaeological 402 visibility of coastal occupation. Both ethnographic and archaeological literature indicate that archaeological 403 sites located more than 5 - 10km from the coast are unlikely to contain substantial evidence for marine 404 resource use. Even distances in the order of only 1 - 2km will dramatically reduce the archaeological density

405 of marine fauna as it becomes mixed with other sources of subsistence (Bailey, 2013:889; Bailey and Flemming, 2008:2155; Bailey and Milner, 2002:5; Bicho et al., 2011; Bird and Bird, 1997; Dusseldorp and 406 407 Langejans 2013; Erlandson, 2001; Fa, 2008; Jeradino, 2016a; Meehan, 1982; Wing, 1977). For example, 408 Meehan (1982) shows that shellfish remains are rarely transported to residential sites >10km from 409 shorelines except in situations where the shells themselves have utilitarian or symbolic value (e.g. Smith and 410 Veth 2004). Bailey and Flemming (2008:2155; Bailey and Milner, 2002) have suggested that the optimum zone for shell midden accumulation is within 1km of the shore, while Dusseldorp and Langejans (2013) 411 412 have documented similar patterns archaeologically for the South African Middle Stone Age. Jeradino 413 (2016a) also shows that shell weight densities drop by half beyond 2km from the South African coast in 414 Holocene deposits. Drawing from these observations, we suggest that coastal sites will contain evidence for marine resource use, but this evidence will become increasingly sparse and dominated by more proximal 415 416 (coastal plain terrestrial) resources with increasing distance from the coastline until evidence for dietary marine resource use drops away to utilitarian or symbolic marine resource use only (as suggested by Meehan, 417 418 1982).

419 For Australia, an average distance within which archaeological evidence for Pleistocene coastal occupation 420 is likely to occur from palaeoshorelines can be calculated using the presence of dietary marine fauna (e.g. 421 non-utilitarian shellfish) as a positive archaeological indicator of coastal occupation. Drawing on the 422 Australian literature for Pleistocene coastal occupation, Table 1 presents this average value calculated based 423 on the distance to a generic palaeoshoreline (using the sea level curve in Ward et al., 2015 paired with 424 bathymetry where possible) at the time when marine fauna is first registered in known Pleistocene deposits. 425 This value can then be compared with the distance of currently available terrestrial sampling points to 426 palaeoshorelines at different points in time. If the latter distance is greater than the former, then it is possible 427 to suggest that evidence for coastal occupations is unlikely to be preserved in terrestrial sampling points. 428 This simple approach avoids problems with marine faunal density (after Jeradino, 2016a) but does not 429 account for other factors like processing costs and marginal returns of other food items (e.g. Codding et 430 al., 2014). Of course, there are significant limitations with this approach including geomorphology and plate tectonics which impact our ability to accurately reconstruct the position of past shorelines using only generic 431 432 sea level curve data and bathymetry (e.g. Larcombe et al. 2018). As such, we suggest the values presented

- 433 in Table 1 be viewed as an approximate guide with an arbitrary error range of ± 2 km. With these limitations
- 434 in mind, our basic approach provides a heuristic but coarse-grained device with which to address the
- 435 question posed above.

Table 1. Tabulated distances to palaeoshorelines for 12 sites at the time when the first marine fauna occurs. Distances to shoreline were calculated using Ward et al.'s (2015) sea level curve and bathymetry in ArcGIS or quoted distances in text. Calibrated dates derive from Ditchfield (2018) and Ditchfield and Morse (in prep) while uncalibrated dates were calibrated using either ShCal20 or Marine20 with local reservoir correction if available. Where possible, dates which derive from the closest stratigraphic or excavation units to the first evidence for marine fauna were selected.

Site	Source	First Marine Fauna (cal. BP)	Distance to Palaeoshoreline (km) 11.3	
Jansz Rockshelter	Przywolnik (2002)	39,140		
C99 Cave	Przywolnik (2002)	37,512	11.8	
Mandu Mandu Creek Rockshelter	Morse (1993a)	36,227	5	
Pilgonaman Creek Rockshelter	Morse (1993a)	12,979	5.1	
Yardie Well Rockshelter	Morse (1993a)	11,536	5.1	
Boodie Cave	Veth et al. (2017c)	48,000	18.6	
Noala Cave	Veth et al. (2007)	13,794	40.4	
John Wayne Country Rockshelter	Ditchfield et al. (2018)	14,932	37.2	
Koolan Shelter 2	O'Connor (1999)	11,577	2	
Bridgewater South Cave	Lourandos (1983)	11,500	3.75	
Koongine Cave	Bird and Frankel (2001)	11,000	15	
Cape Duquesne Middens	Richards (2012)	11,100	1.3	
Average		-	13.1	

⁴⁴¹

The heuristic analysis of Australian Pleistocene sites suggests that most evidence for marine resource use 442 443 will accumulate within 13.1 \pm 2 km of shorelines with increasing abundance towards it. A review of 444 bathymetry shows that, in most instances, the current Australian shoreline is more than 13.1 ± 2 km distant from palaeoshorelines until the terminal Pleistocene (Figure 1). Figure 1 shows the Australian terminal 445 446 Pleistocene shoreline at 12,000–13,000 cal. BP along with the possible palaeoshoreline at 13.1 ± 2 km from the current Australian coastline. It shows that only four major areas should register coastal archaeological 447 448 records at this time: northwest Australia (especially Cape Range), some areas of South Australia, Tasmania and a portion of the New South Wales coast (see Williams et al., 2018 for comparable results). Before this 449 450 time, sea levels were always lower. While we know there is a record of coastal archaeology in the northwest, 451 there are only a few Pleistocene coastal records for the other areas and none that date to before 13,000 cal. BP. This likely results from a lack of targeted sampling and substrate type (e.g. alluvial plains on the New 452 South Wales coast and sea cliffs in South Australia and Victoria; see Bird, 2008). Although, given that coasts 453 are dynamic, not all coasts were productive during the Pleistocene in Australia, meaning that it is possible 454 455 that the lack of Pleistocene coastal records for parts of the four identified regions may be a real 456 phenomenon (as argued by Bowdler 2010 for much of the east coast).

Overall, however, this basic analysis suggests most Pleistocene (especially dense) evidence for marine 457 resource use will be drowned especially before 12,000 - 13,000 cal. BP. As such, it should come as little 458 surprise that almost no Pleistocene shell middens (e.g. greater than 50% shell matrix) are known on 459 460 contemporary shorelines and that, where Pleistocene marine faunal remains are present (e.g. Cape Range), they are sparse and mixed with terrestrial fauna. For example, in southwest Victoria, both Bridgewater 461 462 South and Koongine Caves register sparse marine faunal assemblages mixed with terrestrial fauna during the Pleistocene - Holocene transition when these locations were between 3 - 15km from the shoreline 463 464 (Bird and Frankel, 2001; Lourandos, 1983; Richards, 2012). By comparison, nearby middens at Cape Duquesne, dating to the same period and being only 1km distant from the shoreline, are dominated by 465 466 economic shellfish (Richards, 2012). This example shows that we should generally expect dense economic 467 marine faunal assemblages to accumulate close to palaeoshorelines. Indeed, it is then no coincidence that Erlandson (2001:321-323; see also Bailey and Flemming, 2008), in his seminal review of coastal archaeology, 468 found that only one trait correlates with the preservation of early coastal archaeological records, and that 469 470 was steep bathymetry (i.e. short distance of the current coastline to palaeoshorelines). Almost no early 471 coastal sites were present adjacent to shallow bathymetry ('procumbent shelves').

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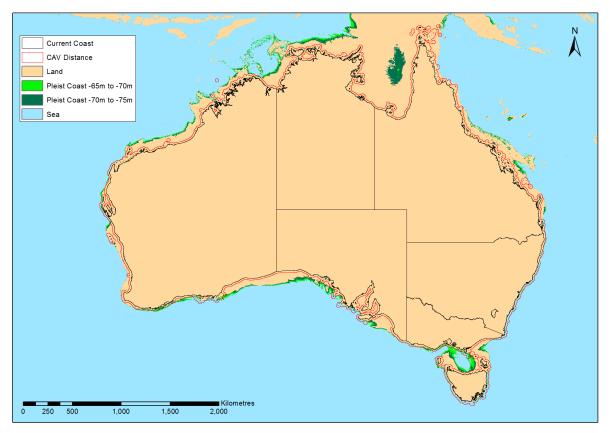




Figure 1. Map showing the Australian continent at 12,000 – 13,000 cal. BP with the possible 13.1km palaeoshoreline and Pleistocene coastal zones within -65 to -75m depth superimposed.

476 Returning to the question posed above then, it is possible to provide a preliminary answer. Unless the current coastline is within approximately 13.1 ± 2 km from palaeoshorelines, there is unlikely to be 477 478 extensive evidence for coastal exploitation in terrestrial contexts in the form of marine resource use, and 479 especially shell middens. Even in steep shelf scenarios (e.g. Cape Range), evidence for significant marine 480 exploitation is unlikely to be dense beyond 1 - 2km from palaeoshorelines (i.e. the 'optimum zone'; Bailey 481 and Flemming, 2008:2155). Rather, the expectation should be that only sparse and mixed marine-terrestrial 482 assemblages should remain from Pleistocene coastal occupations at least until the terminal Pleistocene to 483 early Holocene. Therefore, it is unlikely that the Australian Pleistocene archaeological record provides a 484 representative account of settlement and occupation since the coastal record is predominantly missing. 485 Morse (1993a) recognised this phenomena some time ago:

The relative paucity of evidence for use of the coast during the late Pleistocene,
compared to that for the Holocene, should not be a surprise. The profusion of shell
middens along the modern Holocene shoreline provides unequivocal evidence that

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the preferred location for the consumption of gathered marine resources such as shellfish is the immediate coastal/littoral periphery (Morse, 1993a:278).

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492 4.2 Coastal Marine Resource Productivity

Having established the likelihood of Pleistocene coastal bias resulting from submergence, it is pertinent to
consider whether past coastlines were productive environments and whether they could support
populations.

496 In Australian research, transgressive regimes have often been associated with higher marine productivity 497 while regressive regimes have been argued to be depauperate (Section 2). O'Connell and Allen's (2012:8, 498 2015:76) model does vary from these simple depictions. Their model posits that sea level transgression on 499 procumbent coasts would be destructive, preventing sedimentary stabilisation and marine productivity due 500 to easily inundated low relief shelves. Although, Veth et al. (2014, 2017b) and Ward et al. (2013, 2014, 2015) 501 have since pointed to robust archaeological evidence for marine productivity along procumbent coasts 502 under post-LGM transgressive regimes such as the presence of Terebralia paulustris remains dating to periods 503 of sea level transgression (see also Barker, 1991, 1999, 2004; Ditchfield et al., 2018; Manne and Veth 2015; 504 O'Connor, 1999 for further archaeological evidence). Numerous macro-tidal estuaries across northern 505 Australia also show significant early-to-mid-Holocene estuarine productivity and sedimentation in 506 conjunction with sea level transgression (e.g. Proske et al., 2014). As Bailey et al. (2015:44) argue, sea level 507 rise brings significant marine benefits including increased nutrient recycling from the seabed to the photic 508 zone (water surface) as productive shallow shelves replace steeply shelving offshore topography. In a similar 509 vein, Woodroffe (1990:488) notes that, given the widespread evidence for resilient early Holocene 510 mangrove communities in northern Australia, large river mouths on Australia's continental shelves must 511 have possessed deltaic and estuarine ecosystems throughout their landward retreat. This corroborates the 512 general idea that transgressive regimes can promote marine productivity, challenging the notion of 513 unproductive procumbent coasts (Beaton, 1995; O'Connell and Allen, 2012). Indeed, the number of sites 514 which show evidence for significant coastal occupation under transgressive sea level conditions is 515 internationally increasing (e.g. Colonese et al., 2011; Dillehay et al. 2017; Goodbred et al., 2020; O'Connor

et al., 2018; Pardo et al., 2016; Prendergast et al., 2016; Reitz et al. 2016; Szabo and Amesbury, 2011; and references therein for each). This burgeoning record strongly suggests that transgressive coasts can be productive in many instances, with the possible exception of very rapid sea level rise in the order of ≥ 6.1 mm year⁻¹ (Saintilan et al. 2020). We conclude here that there is little basis to assume that Pleistocene coasts were categorically low in marine productivity during transgressive regimes.

521 This leaves us to consider the effects of regressive regimes, and especially during glacial periods, on coastal 522 productivity. In contrast to assumptions of regressive marine zones being unproductive (Section 2.2), Ward 523 et al. (2013, 2015) have emphasised that periods of regression can produce broader coastal plains which 524 will generally propagate larger tidal ranges and create greater opportunity for sediment accumulation, both 525 of which can enhance marine productivity. For example, Fa's (2008) work demonstrates that broader coasts 526 are characterised by larger tidal amplitudes which can support large intertidal molluscan biomasses that are 527 capable of withstanding significant intertidal exploitation. However, it is worth noting that palaeotides were 528 not the same as those observed today (Haigh et al., 2019), with changes in sea level effecting tidal resonance on the continental shelf (e.g. in the Timor Sea). Recent exploration of the effects of changing sea level on 529 530 tidal range and currents by Kuijjer et al (2022) indicates lower tidal regimes (4m tidal range as opposed to 531 present day scenario of 5.65m) along the northern Australian shelf and Bonaparte Gulf in MIS 4 regressive 532 phases. This indicates that whilst a broad 'procumbent' coast was exposed, tidal range and therefore the 533 intertidal zone was not as great as today in this region, whilst other regions had greater tidal regimes.

534 In addition, sea level regression - transgression can create a greater total continental shoreline meaning that, 535 should some marine zones be productive, there may be greater total opportunity to exploit them. For 536 example, the changing shorelines of the modern Great Barrier Reef region across the Holocene 537 transgression can be considered through quantifying sinuosity, a measure of the departure of the shape of 538 the coastline from a straight line. Figure 2 shows the coastline of northeast Australia at approximately modern levels (8,000 BP) and lowstand (≥15,000 BP), with an intermediate figure at -50m showing 539 540 thousands of islands created around last interglacial reef and landscape features. This, in theory, would not only increase length of coastlines but also the extent of intertidal zones and, ultimately, marine productivity. 541

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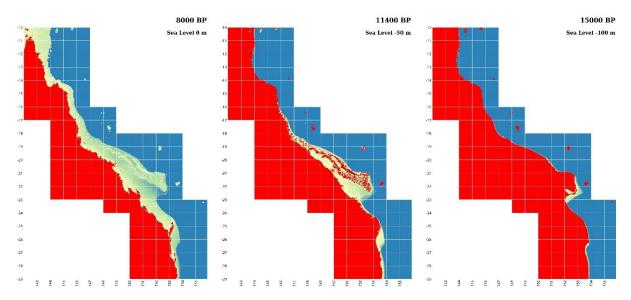


Figure 2. Queensland coastline shown at selected time-slices. Red indicates extent of terrestrial coasts and islands. Note that islands
 identified in the 11,400 BP time-slice include topographic features from Holocene reef-growth. However, the general pattern
 reflects landforms created during the last interglacial period.

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While these ideas may support regressive marine productivity in theory, direct evidence for marine
productivity on regressive shorelines is required. Following Chappell's (1993; Chappell and Thom, 1977)
suggestions, marine productivity can be characterised by reasonably high sedimentation in the near-shore
zone in association with dietary marine fauna. Recent research addresses these criteria.

552 Ishiwa et al. (2016) recently obtained an offshore core from the last glacial coastline of the Bonaparte Basin. Shell fragments appear from 29,000 cal. BP, including Anadara sp., indicating estuarine conditions as sea 553 554 level receded towards the Bonaparte coast. On the same Bonaparte Basin glacial coast, De Deckker and Yokoyama (2009) and Fogg et al (2020) track coastal evolution under regressive conditions where open 555 marine conditions grade into shallow marine, marginal marine and then to estuarine brackish conditions. 556 Fogg et al (2020) use industry collected 3D seismic data to map shoreline position along with changing 557 558 littoral and estuarine features in the Bonaparte Gulf. Preliminary results support the available core data, indicating the now submerged shelf hosted palaeofluvial systems, estuarine areas and regions with potential 559 for mangrove growth in low stand periods. Using the core data, De Deckker and Yokoyama (2009:87) note 560 561 that dense molluscan remains accumulated under shallow water conditions including bivalves, echinoid 562 spines and scaphopods, while, in the brackish sediments are insect remains and vegetal fibres which strongly suggests heavy continental freshwater influence. De Deckker and Yokoyama (2009) even recovered fish 563

fragments from brackish, estuarine units during fully glacial periods. The density of shellfish and the 564 thickness of shell layers also demonstrates rapid sedimentation during this time. The Bonaparte core records 565 566 definitively show that productive marine conditions were associated with regressing shorelines and that 567 they remained productive throughout the Last Glacial Maximum. Although this is only one example from 568 one coastline, it is critical because it shows that, at the very least, regressive marine zones were not 569 universally unproductive. Beyond the Bonaparte Gulf, modelling by Hinestorsa et al. (2016) suggests that 570 sediment accumulation may have been significant during both regressive and transgressive regimes at the 571 Noggin Passage, potentially leading to productive estuaries and mangrove systems on parts the Pleistocene Great Barrier Reef (GBR) coast. Webster et al. (2018), using drill cores, also show that the GBR was resilient 572 573 to both regressive and transgressive regimes, never being completely 'decimated' by sea level fluctuation. 574 In Africa, there are also numerous examples of continued coastal occupations during Pleistocene regressive 575 conditions (e.g. Jeradino 2016a, 2016b; Loftus et al., 2019; Niang et al., 2018; Will et al., 2016, 2019; see 576 also Zilhao et al., 2020).

577 As such, we conclude that current models for Pleistocene marine productivity in Australia may under-578 estimate the resilience of marine ecosystems in response to sea level fluctuations (Erlandson, 2001, 2012; 579 see also Barker, 1999, 2004). While we already have some evidence for coastal occupation on regressive and 580 transgressive coasts from northwestern Australia (see above), there is no reason other coasts across 581 Australia could not be intensively occupied where marine resources were systematically exploited. Indeed, 582 coasts may very well have served as refugia during the LGM (Morse 1993), a subject of further planned 583 work at Cape Range in northwest Australia by the authors. Internationally, there are also examples of coastal 584 occupation during glacial conditions. For example, based on oxygen isotopic analyses, Prendergast et al. 585 (2016) have shown that late glacial coasts were occupied year-round at Haua Fteah in Libya. Such analyses 586 may yield similar results in Australia.

587

588 4.3 Coastal 'Terrestrial' Productivity

589 One of the outstanding issues in considering Australian coastal productivity is that most research only590 considers the marine zone when assessing productivity (e.g. Section 3; but see Hallam, 1987). However,

researchers are beginning to illustrate that coastal economies will also include suites of terrestrial coastal plains fauna and flora (Bailey et al., 2008; Ditchfield et al., 2018; Erlandson and Braje, 2015:34; Hallam, 1987; Klein and Bird, 2016). The coastal plain is an important component of the coast for mobile coastal hunter-gatherers. After-all, coastal hunter-gatherers do not live in the intertidal zone, they occupy the coastal plain and so should be expected to use its resources. Essentially, a coastal economy occurs *from the land* and the interoperability of the two zones is theoretically and practically unified (see Roe 2000).

597 In this context, it is significant that most Pleistocene coasts fronted much larger coastal plains than at 598 present and there is good reason to believe that at least some were very productive. Faure et al. (2002) have 599 observed that emergent Pleistocene coastal plains were likely well-watered environments, even dubbing 600 them as 'coastal oases'. This is because sea water columns exert significant pressure on continental ground-601 water reservoirs when the water column lies over the continental shelf (high sea levels). Under conditions 602 of lower sea level, when the water column pressure is removed from the continental shelf, groundwater can 603 become emergent across the shelf in the form of springs, soaks and streams. As Faure et al. (2002:52-53) note, a 120m drop in sea level is equivalent to raising the continental water table by 120m. This effect will 604 605 be particularly profound on large, wide, low relief shelves (of which Australia has many; Yokoyama et al., 606 2001) where significant water columns are resident during high sea levels. Indeed, the northeast Australian 607 coastline likely hosted large freshwater lagoon systems before sea-level rise breached the continental shelf, 608 with karstified hills along the continental margin from previous reef building episodes during high stands 609 blocking the outlet of rivers to the shelf edge (Woolfe et al., 1998; Dunbar and Dickens, 2003). Benjamin 610 et al. (2020) have also located evidence for the availability of freshwater on the Pleistocene coast in the 611 form of drowned freshwater springs near Cape Bruguieres, Murujuga.

Under these conditions, sources of coastal freshwater were likely available throughout the Pleistocene. Indeed, during arid glacial conditions (when ground-water accessibility may have been, theoretically, at its greatest), the coast possibly represented a very significant refugium for terrestrial flora, fauna and people even in the absence of productive marine ecosystems. Bailey et al. (2007) have made similar suggestions for Arabia while Dusseldorp and Langejans (2013) have suggested that some coastal plain sites may have been preferentially occupied to maximise terrestrial coastal resource exploitation. Manne (1998) proposed that a rich range of resources became available during the Pleistocene-Holocene transition in the area of the 619 Montebello Islands, as groundwater levels relating to changes in sea level increased the availability of near-620 surface freshwater, which in turn increased productivity in the coastal hinterlands. Ditchfield et al. (2018) 621 have suggested that both marine and coastal plain ecosystems were economically integrated during the 622 terminal Pleistocene on Barrow Island. Indeed, perhaps coastal plain resources were the main drawcard for 623 some Pleistocene coastal hunter-gatherers where any settlement strategy capable of effectively exploiting 624 both terrestrial and inter-tidal zones would carry a significant advantage (Bailey and King, 2011:19-20; see 625 also Kyriacou et al., 2014). Interestingly at Gua Makpan on Alor Island (Indonesia), and other sites from 626 isolated Wallacean islands, Roberts et al (2020; see also Kealy et al., 2020) used stable isotope analyses of 627 human teeth to reconstruct the diet of past cave occupants during the Pleistocene and Holocene. The 628 archaeological record at Gua Makpan indicates a focus on marine subsistence during occupation which began at 40,208 – 38,454 cal. BP. Initial reliance on urchins and barnacles expanded to a wide range of 629 630 gastropods and bivalves as well as medium-sized inshore fish and Scombridae (mackerel, bonito and tuna). Terrestrial macrovertebrates are only present as a minor component throughout the sequence and includes 631 giant rats, fruit bats and turtles (Kealy et al., 2020). However, the isotopic analysis from Gua Makpan and 632 other sites shows the diet includes a broader use of terrestrial interior resources, particularly after 20,000 633 cal. BP (Roberts et al., 2020). If this analysis is correct, it would indicate that, even on islands adjacent to 634 635 steep shelf topography, coastal occupation did not solely focus on marine foods but actively incorporated 636 accessible terrestrial components which may also include plant resources. As Roberts et al. (2020:8) suggest, 637 the stable isotopic data highlights the need to pay further attention to the potential contribution of plant 638 and terrestrial animal resources to human diets on tropical islands and, as we might suggest, to coastal 639 settings more generally.

In future discussions it may be useful to differentiate between coastal and marine adaptations where the former is characterised by the use of both coastal terrestrial and marine resources while the latter is characterised by systematic use of marine resources (cf. Jeradino, 2016b). One way to test this proposition is through detailed examination of archaeological faunal and floral assemblages (see below). Overall, we suggest that the archaeological record of coastal plains will be just as critical as sites immediately adjacent to the shore for reconstructing past coastal adaptations (Erlandson, 2015; Erlandson et al., 2011; Jones, 1977).

647 5. Future Directions

We have argued that most of the evidence for Pleistocene coastal occupation in Australia is now submerged, 648 which logically explains why there is so little recoverable evidence and, when it does occur, it is adjacent to 649 650 'precipitous' coasts. Based on available literature in association with archaeological and 651 palaeoenvironmental evidence, we have suggested that there is no reason to assume that coastal landscapes 652 were unproductive or unoccupied. This discussion provides a framework for conceptualising the 653 occupation of Pleistocene coasts that are currently drowned and largely inaccessible for sampling purposes. 654 It also provides a platform with which to consider future research and what additional evidence might be 655 required to develop our understanding of Australian Pleistocene coasts. As with some other global study 656 areas, we might now consider how the systematic use of coastal resources on productive Pleistocene coasts 657 would be expressed archaeologically over space and time.

We identify three major areas for future research: faunal studies, stone artefact research and coastal
modelling. Of course, there are further areas of future work including geomorphology (e.g. Jeradino 2016a;
Ward et al. 2017, 2018) and underwater archaeology on submerged sites (e.g. Benjamin et al. 2020; Benjamin
and Ulm 2021; McCarthey et al. 2022) but we wish to focus on these three here.

662

663 5.1 Faunal Research

So how might we move forward, with much of our Pleistocene coastlines lying submerged beneath the sea? 664 One way is to further explore the nature of diet breadth in coastal sites. While few sites exist for 665 interrogating coastal faunal records of the late Pleistocene to early Holocene in northern Australia, some of 666 667 the best evidence for coastal living is in the Montebello-Barrow Islands complex. Records from Noala, 668 Haynes and Boodie Caves suggest that as the coastline progressively came closer to these sites, people took advantage of an increasingly diverse range of resources from a variety of marine and terrestrial habitats 669 (Manne and Veth, 2015; Veth et al., 2017; Veth et al., 2007). Both the greatest intensity of site use and 670 671 diversity of taxa recovered, occur when the coastline was within 0-10km of these sites (Manne and Veth, 2015; Veth et al., 2017; Veth et al., 2007). 672

If viewed from the perspective of evolutionary ecology, broad diets are traditionally associated with a 673 674 lowered abundance of higher ranked resources, with diet diversification following a reduction in encounter-675 rates with highly ranked resources (Stephens and Krebs, 1986). The broad diet observed in the Montebello 676 and Barrow Island sites could thus be interpreted as being driven by a limited abundance of higher-ranked prey, compelling people to include a broad selection of prey items in their diet. However, in northwest 677 678 Australia there is little evidence for narrow diets from even the earliest phases of occupation and instead, 679 early sites suggest that a broad diet was the norm. For example, interior sites such as Riwi and Carpenters 680 Gap demonstrate that a diverse suite of taxa were targeted, from earliest settlement onward (Balme et al., 2019; Maloney et al., 2018). In the northwest of Australia, a broad diet may not be indicative of a response 681 682 to a reduction in higher-ranked prey, but instead, signify the most parsimonious diet for a region with a 683 diversity of small to medium bodied game coupled with either sporadic or highly seasonal precipitation. Perhaps then, coastlines may have always been particularly attractive due to the additional resources 684 afforded from the combination of marine and coastal hinterland habitats. 685

Recent work by Dillehay et al. (2017) on sites on the Chicama coast of northern Peru lends support for the 686 687 view that coastal environments were appealing in terms of their rich diversity of resources. At 15,000 cal. 688 BP, the coastline was approximately 30km to the west of the current Chicama coast and by 10,000 cal. BP, the coast was 8 - 10km away (Dillehay et al., 2017). However, already 15,000 years ago, people in this region 689 690 were exploiting shoreline, estuarine and hinterland resources in the form of both fauna and flora while also 691 indicating a familiarity with more interior resources. Dillehay et al. (2017) argue that the abundance of 692 resources provided by the intersection of different ecological zones in this region encouraged coastal 693 settlement and consequently slowed migration into the interior.

Turning back to northwest Australia, we argue for a renewed focus on the Cape Range-Barrow-Montebello region. This is an area known for its unique, well-preserved archaeological and faunal record and has excellent potential for improving our understanding of how people utilised and moved between the diverse sets of resources that likely occurred between the hinterland and the coast. Our understanding of what diet breadth over time actually means in this region would be improved by examining larger sample sizes of fauna. By comparing multiple faunal assemblages from across this region, issues of sample size – which so often plague our understandings of past economies in northern Australia – may be alleviated, allowing for
a more detailed and nuanced understanding of the past 50,000 years.

702

703 5.2 Stone Artefact Research

704 There is very little research on stone artefact assemblage formation processes or technology on Pleistocene 705 coasts, which makes this an important research area. How do the technologies and assemblage patterns on 706 Pleistocene coasts compare with those in other terrestrial environments? If coasts were regularly occupied during the Pleistocene, technological patterns may differ significantly from better documented inland areas. 707 708 Since stone artefacts also preserve well in the archaeological record, they will be critical for identifying 709 drowned archaeological sites (Stanford et al., 2014). Stone artefact assemblages can also be used to 710 reconstruct important regional-scale behavioural processes such as mobility. We know little about 711 Pleistocene coastal mobility, though occupation is currently thought to be short-term (O'Connell et al. 712 2012). The mobility of past people is also linked with environmental productivity where certain mobility 713 patterns are linked to the spatial and temporal structure of resources throughout a landscape (e.g. Binford, 1980, 1982; Brantingham, 2006; Grove, 2009; Hamilton et al., 2016; Kelly, 2013). In this way, targeted 714 research on available assemblages can help reconstruct little known mobility patterns, which will also help 715 716 inform on Pleistocene coastal productivity.

717 One of the key areas that is worth highlighting is sourcing the stone artefacts discarded across Pleistocene 718 coasts. Sourcing provides one measure for the distance and direction of human movement, and connects 719 the point of discard with the source location (e.g. quarried outcrop or secondary deposit; e.g. Ditchfield et 720 al. 2021). This may have significant potential for Pleistocene coasts since sourcing non-local stone artefacts 721 may link interior landscapes with the coast, possibly demonstrating trade networks or seasonal movements. 722 Ditchfield and Ward (2019) have begun sourcing work using Pleistocene coastal assemblages from Cape 723 Range and Barrow Island. While their goal was to establish the source locations for locally derived 724 lithologies, this research also established which raw materials were non-local. For example, on Barrow 725 Island, both stratified and open assemblages contain non-local lithologies that derive from igneous or 726 siliceous sedimentary source locations. None of these occur on the island so must derive from the presentday mainland or now-drowned source locations. Further work will help pinpoint the source locations forthese non-local materials.

729

730 5.3 Coastal Modelling and Submerged Archaeology

Coastal modelling can be a key tool in helping to generate an understanding of past occupation on 731 732 Pleistocene coasts and, to date, has taken many forms. For example, predictive models can be used to 733 inform on the potential locations that submerged archaeological sites might preserve. The modelling work 734 undertaken by Veth et al. (2019) for Murujuga provides an example of predictive modelling. Veth et al. 735 (2019) argued for a multi-staged assessment strategy for examining the possible continuity of terrestrial and 736 maritime archaeological landscapes. Terrestrial site patterning through time was based on a large sample of 737 dated middens and occupation sites, considering marine processes, and surviving contexts. Based on this 738 terrestrial dataset, targeted marine survey can be deployed at various scales using remote sensing, diving, 739 and coring (Benjamin et al., 2018). The predictive model concluded there were five likely targets for 740 submerged sites ranging from well stratified and structured occupation sites including middens through to lag deposits of artefacts and stone features on hardpan. Grøn et al (2022) have recently critiqued how the 741 application of analogue topographical models, along with their assumptions of productivity, resource 742 743 distribution through time and inundation, will likely be far more complex than continuity models and thus 744 both the accurate location of submerged sites and their dating in Australia are in their infancy (e.g. Benjamin 745 et al. 2020)

746 Important components of shorelines can also be established using modelling techniques based on industry 747 derived seismic data. O'Leary et al. (2020) have done this for the North West Shelf around Barrow Island 748 and the Montebello Islands. Their analysis has identified features of a past stable coastline including dune 749 systems likely dating to MIS3 (57 - 29k). This shows that sea levels were stable enough to produce 750 prominent features for a period during MIS3 and that this coastal landscape included tidal flats, estuarine 751 channels, coastal lagoons and a coastal plain some 15km wide. O'Leary et al. (2020) argue that these 752 environments likely produced a productive marine ecosystem including mangrove communities like the 753 present-day Pilbara coastline. High-resolution palaeogeographic reconstructions have also been undertaken 754 using 2D and 3D seismic data in the Bonaparte Gulf (Fogg et al., 2020). Seismic interpretation alongside 755 extant core data allows the mapping of changing coastal configurations, including palaeochannels, estuaries, 756 headlands, and offshore islands (Fogg et al., 2020). This can form the basis for mapping the potential 757 locations and preservation states of submerged sites too. Sub-bottom profiling is another critical step for future modelling and palaeoshoreline mapping since it corrects for the impact of post-transgressive 758 759 sediments on offshore platforms (Erlandson, 2021). Modelling of the maritime environment and 760 palaeolandscape to include ocean circulation models and Pleistocene tidal models (for instance, Kuijjer et 761 al., 2022) will also allow for a better in-depth understanding of maritime and coastal conditions, which will 762 aid understanding of coastal productivity and maritime mobility.

763 Locating evidence for submerged archaeology on the continental shelf is clearly an important next step. 764 This effort is well underway, Benjamin et al. (2020) have recently reported the first submerged Aboriginal 765 archaeological site with >200 stone artefacts recorded at Cape Bruguieres located off the Murujuga 766 coastline. Minimum dates place site occupation before 7000 cal. BP. While there are many difficulties with 767 locating and investigating drowned sites in Australia (see Benjamin et al., 2020; Benjamin and Ulm 2021; 768 McCarthy et al. 2022; Veth et al. 2020), we believe this will be the first of many Aboriginal archaeological 769 sites located in drowned marine contexts which, ultimately, adds credence to our suggestions that most of 770 the Australian Pleistocene coastal archaeological record is drowned (Section 4.1).

771

772 6 Conclusion and Final Remarks

773 Although numerous spatio-temporal processes condition the exact productivity of any given coastline, 774 Pleistocene coasts can be conceptualised as potentially productive but taphonomically biased (through sea 775 level rise and associated processes). The discussions and arguments put forward in this paper provide a 776 basic framework, and a different set of environmental expectations, within which to assess independently 777 reconstructed occupation patterns to those proffered by other models for Pleistocene coastal occupation 778 in Australia (e.g. Beaton, 1995; O'Connell and Allen, 2012). We believe that the case for Pleistocene coastal 779 occupation and coastal (marine and terrestrial) productivity is high. We do not believe that stable coasts 780 (without fluctuating sea levels) are a prerequisite for productivity. Instead, we believe that the lack of

781 archaeological evidence for coastal occupation and marine resource exploitation in Australia is mostly due to a drowned record lying on continental shelves with only terrestrial-marine mixed assemblages persisting 782 in areas of relatively steep continental relief. This may seem like an obvious conclusion to some, but our 783 784 review of the literature has shown that this is not uncontested. Regardless, while most of the record for previous coastal occupation lies submerged on continental shelves, the behaviour of people, such as 785 786 mobility across these extensive landscapes, remains poorly understood. In order to recalibrate the 787 terrestrially biased archaeological records to be more representative and inclusive of Pleistocene coastal 788 occupational patterns, it is now crucial that more research is undertaken targeting landscapes that can provide Pleistocene coastal windows, such as continental islands and uplands once close to the shelf such 789 790 as Cape Range. A new generation of research is planned for both these areas and for submerged records 791 off the southern and northeast tropical coasts.

792

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