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1 **Rainfall and its possible hysteresis effect on the proportional cover of tropical tidal-**  
2 **wetland mangroves and saltmarsh–saltpans**

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11 **Abstract.** Mangrove–saltmarsh tidal wetlands are highly dynamic ecosystems, responding and adapting to  
12 climate and physical conditions at all spatial and temporal scales. Knowledge of the large-scale ecosystem  
13 processes involved and how they might be influenced by climate variables is highly relevant today. For tidal-  
14 wetland sites well within the latitudinal range of the mostly tropical mangrove communities, we confirm that  
15 average annual rainfall influences vegetative cover, as well as species composition and biomass of tidal  
16 wetlands. On the basis of 205 largely unmodified, tropical and subtropical estuaries of northern Australia, a  
17 sigmoidal relationship, with a centroid inflection point ~1368 mm, was derived between rainfall and the relative  
18 amounts of high-biomass mangroves and low-biomass saltmarsh–saltpan vegetation. The presence and  
19 probability of observed combinations of these community types were quantified using the wetland cover index,  
20 **which is** the ratio of total mangrove area to that of mangroves plus intertidal saltmarsh and saltpans.  
21 Accordingly, periodic changes in rainfall trends are likely manifest as either encroachment or dieback of  
22 mangroves along the ecotones separating them from tidal saltmarsh–saltpans. Presented is a new conceptual  
23 framework and model that describes how such ecosystem-scale processes take place in tropical and subtropical  
24 tidal wetlands.

25 **Summary.** A revised look at tidal wetlands in remote northern Australia has uncovered a startling new  
26 understanding of their responses to changing climate. Mangrove and saltmarsh–saltpan communities appear to  
27 act as one combined ecological niche. As rainfall conditions vary, one seems to expand while the other contracts,  
28 with no net change in total area. Despite all the pressures on tidal wetlands, it seems that moisture and salinity  
29 are the dominant and predictable determinants of relative cover for these globally recognised habitats of tropic  
30 and subtropic regions.

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33 Ecotone shift in tropical tidal wetlands

34 **Additional keywords** climate change, ecosystem state change, ecotone shift, tidal saltmarsh, wetland cover  
35 index.

## 36 **Introduction**

37 Mangroves and saltmarsh tidal wetlands occur widely within climate and physical conditions that  
38 vary at broad spatial and temporal scales (Duke *et al.* 1998; Saintilan 2004; Feller *et al.* 2010). The  
39 presence of mangrove forests today, after more than 50 million years (Duke 2017), is testament to  
40 their capabilities for longer-term survival and adaptive success. However, these inherent capabilities  
41 do not necessarily ensure their continued survival in the face of growing and cumulative direct human  
42 influences such as pollutants and physical damage, coupled with rapid changes in climate, sea-level  
43 rise and other detrimental pressures (Duke 2014; Spalding *et al.* 2014; Harris *et al.* 2018). Recent  
44 events, such as the widespread mass dieback of mangroves in Australia's Gulf of Carpentaria (Duke *et*  
45 *al.* 2017), provide a cogent indication of the limits and vulnerabilities of these otherwise robust and  
46 durable ecosystems. At this time, there is great relevance in explaining the role of large-scale  
47 processes that are likely to bring change and may, or may not, be linked to global climate change  
48 (IPCC 2014). Where the extent and character of tidal-wetland vegetation are constrained by climatic  
49 factors (cs. Duke *et al.* 1998), the resulting changes might subsequently affect function as well as the  
50 many ecosystem services derived from these habitats, such as their carbon-storage potential (Duke *et*  
51 *al.* 2007; Wilkie 2007; Ward *et al.* 2016).

52 Several physical and climate drivers are recognised for their ability to limit and influence processes  
53 within coastal–estuarine wetland habitats (Duke *et al.* 1998). These specifically include rising sea  
54 levels leading to upland encroachment (Jiang *et al.* 2015; Lovelock *et al.* 2015), increases in  
55 temperature, driving species shifts to higher latitudes (Osland *et al.* 2013; Saintilan *et al.* 2014), and  
56 variations in rainfall, affecting vegetation cover within the tidal-zone habitat (Eslami-Andargoli *et al.*  
57 2013; Osland *et al.* 2014, 2016). Although these studies and others have made substantive  
58 contributions, fundamental gaps remain in our understanding of how tidal-wetland ecosystems  
59 respond to change, and how particular drivers specifically influence ecosystem processes.  
60 Furthermore, both ecological knowledge and the underlying physical data appear lacking. In a recent  
61 review of mangrove cover from around the world, Friess and Webb (2014) reported that the  
62 fundamental parameter of spatial extent was deficient, leaving year-to-year quantification of longer-  
63 term trends fraught at best (also see Jennerjahn *et al.* 2016; Rogers *et al.* 2016). This situation appears  
64 compounded further by a limited appreciation of how landscape-scale ecological processes might be  
65 affected as specific climate variables change.

66 In the present assessment, we focused first on temperature and then rainfall as the two dominant  
67 climate drivers. Both are known to influence the relative cover of mangrove and the extent of  
68 saltmarsh tidal wetland, as defined in the following two groupings:

69 (1) *Temperature-influenced changes* to areas of mangrove v. saltmarsh vegetation in the vicinity of  
70 higher-latitude limits of mangrove. These studies include mangrove expansion linked to periods of  
71 greater warmth (Gilman *et al.* 2008; Osland *et al.* 2013, 2014, 2016; Saintilan *et al.* 2014) and the  
72 alternate scenario, i.e. mangrove dieback or retreat linked to severe frosts and cooling (Kao *et al.*  
73 2004; Ross *et al.* 2009; Feller *et al.* 2010; Saintilan *et al.* 2014).

74 (2) *Rainfall-influenced changes* to mangrove v. saltmarsh vegetation within the latitudinal limits of  
75 mangroves, where mangrove expansion was linked primarily to wetter conditions (Duke *et al.*  
76 2003; Gilman *et al.* 2008; Eslami-Andargoli *et al.* 2013; Duke 2014). The alternate scenario is that  
77 mangrove loss or retreat is linked to drought and decreased precipitation (Duke *et al.* 2003;  
78 Gilman *et al.* 2008; Duke 2014; Duke *et al.* 2017). A further subgroup concerns rainfall-  
79 dominated saltmarsh dieback at sites beyond the higher latitudinal limits of mangrove  
80 occurrences. In these cases, there may have been links to drought, decreasing rainfall and sea-level  
81 changes (McKee *et al.* 2004; Silliman *et al.* 2005).

82 Each of these instances featured direct links between temperature and/or rainfall, and the ecological  
83 responses of tidal wetland vegetation as physical conditions changed. Although such responses are  
84 likely to be confounded at higher-latitude limits of mangroves, it seems reasonable to assume for  
85 locations well within the latitudinal range of mangroves, that temperature might be less influential.  
86 The observed variability in the vegetative extent among the low-latitude sites appears consistent with  
87 the greater importance of rainfall in influencing both the biodiversity and biomass of mangrove  
88 vegetation (Duke *et al.* 1998), as well as the percentage cover.

89 Finding a definitive relationship between vegetation cover and climate variables, particularly  
90 between rainfall and observed tidal-wetland vegetation states, has proven elusive. This is despite  
91 observations made more than half a century ago by an eminent field botanist, Fosberg (1961), who  
92 reported that ‘vegetation-free zones’ of coastal tidal wetlands of northern Australia and Central  
93 America were more expansive in locations with drier climates. Fosberg (1961) went on to propose a  
94 direct relationship between the vegetation cover of tidal wetlands and rainfall (exemplified in Fig. 1).  
95 Several studies since have considered this proposal (Bucher and Saenger 1994; Danaher 1995; Digby  
96 *et al.* 1999; Mount and Bricher 2008; Gilman *et al.* 2008; Eslami-Andargoli *et al.* 2013).

97 Eslami-Andargoli *et al.* (2013) analysed changes in the areal extent of mangrove and saltmarsh  
98 within Moreton Bay in Queensland, Australia, and found that mangrove expansion into saltmarsh-  
99 saltpan areas was linked to rainfall. Importantly, this was indicative also of how these two habitats  
100 could be influenced by each other. However, the strength of the relationship was considered weak,  
101 because rainfall explained less than 51% of the variance. They concluded that although additional  
102 factors such as localised sea-level variability may have influenced the relationship, it was more likely  
103 that there were serious confounding effects from surrounding anthropogenic influences, especially  
104 where these differed greatly between estuaries. More recently, a non-linear relationship between

105 mangrove vegetation cover and rainfall was reported by Osland *et al.* (2014, 2016) for tidal wetlands  
106 of the Gulf of Mexico, USA. Although these studies appear to have direct relevance to our inquiry, it  
107 is likely that their results were confounded by the influence of low temperature because mangrove  
108 sites were mostly at, or close to, their upper-latitude (= low-temperature) limit.

109 Bucher and Saenger (1994) used data from both temperate and tropical locations in Australia and  
110 described similar relationships of wetlands with rainfall and temperature. Wetland cover was  
111 expressed as a proportion of the sum of estuarine areas by both intertidal and subtidal habitats  
112 (mangroves, saltmarsh–clay pan, seagrass, lower intertidal flats and open water; Bucher and Saenger  
113 1994). The inclusion of, arguably, independent subtidal habitats in the index, coupled with likely  
114 inherent-scale errors in their spatial quantification (e.g. the inclusion of below-water seagrass habitat)  
115 undermined the reliability of Bucher and Saenger’s index. Bucher and Saenger (1994) reported linear  
116 trends in mangrove and saltmarsh cover with rainfall, but made no comment on a possible co-  
117 relationship between the vegetation types. Here, we present an extended assessment of how wetland  
118 cover relates to temperature and rainfall, using a specifically targeted index, the wetlands cover index  
119 (WCI). Fig. 2 illustrates how the choice of index affects estimates of vegetation cover.

120 This study aimed to identify the likely key ecosystem processes involved in governing wetland  
121 cover, and how these might influence the temporal change at individual locations. The following three  
122 hypotheses were tested: (1) linear or non-linear relationships exist between vegetative cover of tidal  
123 wetlands and rainfall; (2) dynamic relationships in the biomass and dominant vegetation states of  
124 tropical tidal wetlands correlate with longer-term trends in rainfall; and (3) the ecotone between  
125 dominant vegetative components shifts in corresponding retreat or expansion of mangroves depending  
126 on longer-term rainfall.

## 127 **Methods**

### 128 *Study system*

129 A key feature of tidal wetlands is their distinctive zonation and uniform vegetation assemblages  
130 (Snedaker 1982; Bunt 1996; Duke *et al.* 1998; Tomlinson 2016). These zones are defined by a small  
131 number of notable ecotones. The two primary ecotones follow either the sea-edge contour at lower  
132 elevations (~mean sea level), or upper-elevation limits for mangroves that generally mark the  
133 landward margin (~highest astronomical tide level). Above the upper ecotone occurs upland, supra-  
134 tidal plants and habitat. Between these two ecotones, there are other distinct zones described by their  
135 often-uniform assemblages of species-specific vegetation units. The most prominent of these is the  
136 ecotone between low-biomass vegetation consisting of herbaceous microphyte-dominated flats of tidal  
137 saltmarsh and saltpans (Adam 1995; Johns 2010; Sainty *et al.* 2012) and the high-biomass  
138 macrophyte-dominated shrub and tree vegetation of mangroves (Duke 1992; Spalding *et al.* 1997;  
139 Tomlinson 2016). An important deduction is that these zones must alter and shift as the constraining

140 factors change, such as sea level, inundation frequency, moisture stress, salinity, temperature and  
141 exposure, along with other direct human influences such as added nutrients.

#### 142 *Spatial data and the study area*

143 Data on estuarine habitats for tropical northern Australian estuaries within Western Australia,  
144 Northern Territory and Queensland were sourced from the OzCoasts database (Australian Government  
145 Geosciences Australia 2017). OzCoasts is maintained by the Australian Government Geoscience  
146 Australia and includes data collated by [Bucher and Saenger \(1994\)](#). These data were supplemented,  
147 for northern Queensland, with survey results from the Cape York Peninsula Land Use Strategy project  
148 ([Danaher 1995](#)). Collectively, these data included 770 estuaries reported for the Australian region.  
149 These data provided classifications of the relative condition of each estuary as pristine, largely  
150 unmodified, modified or extensively modified. The analysis in the present study was restricted to 205  
151 pristine or largely unmodified estuaries in tropical Australia, of a scale of 10–100 km<sup>2</sup> ([Table 1](#)).  
152 Estuaries <10 km<sup>2</sup> were excluded to minimise mapping error. Ozcoasts mapping data followed the  
153 National Intertidal–Subtidal Benthic (NISB) Habitat Classification Scheme ([Mount and Bricher 2008](#)),  
154 which is more accurate in larger systems owing to better resolution of available aerial photography  
155 and lessened scale-dependent mapping errors. Estuaries >100 km<sup>2</sup> were removed to minimise the  
156 impact of within-estuary heterogeneity of rainfall. Measures were taken at any spatial scale, but they  
157 represented significant portions of any estuarine system, having most upper-tidal levels represented.  
158 The sites sampled (situated between 11 and 23°S) had maximum temperatures between 30 and  
159 36°C and minimum temperatures between 18 and 24°C, on the basis of the 20–30-year averages.  
160 Accordingly, changes in temperature were not considered to have a major defining influence on the  
161 sites considered.

#### 162 *Rainfall data*

163 Rainfall data were sourced from the Australian Government Bureau of Meteorology (weather  
164 stations, and maps showed isohyets of rainfall zones (Australian Government Bureau of Meteorology,  
165 2017). Rainfall was quantified as the average annual rainfall (AAR) received by each estuary and  
166 catchment over the previous 20–30 years. For estuaries >100 km<sup>2</sup>, the AAR data available from the  
167 Australian Bureau of Meteorology was generated not only on a time-based statistic, but also on the  
168 result of an average of all rainfall data available over each catchment.

#### 169 *Wetland cover index and the notable vegetation ecotone*

170 The dramatic and often stark height and biomass gradient between woody plant communities  
171 ([tree/shrub](#)) and [saltmarsh-saltpan](#) plants forms a clearly defined and identifiable ecotone boundary, or  
172 zone limit, within tidal wetlands. This ecotone marks a distinct limit of mangrove establishment along  
173 the tidal elevation profile defined by seedling desiccation and low soil water content ([cs. Clarke and](#)  
174 [Myerscough 1993](#)). The WCI was applied to broadly quantify the position of this natural ecotone, on

175 the basis of the relative abundance of mangrove vegetation within the tidal wetland overall, and  
176 specifically, as the ratio ‘mangrove area : total area of tidal wetland’, where ‘mangrove’ equals  
177 mangrove-dominated, ‘tidal wetland’ equals mangrove plus tidal saltmarsh and saltpan, ‘saltmarsh and  
178 saltpan’ equals tidal saltmarsh plus high intertidal saltpans. The WCI provided a simple measure of the  
179 dynamic state of tidal-wetland vegetation cover for all 205 estuarine catchments (Table 1).

#### 180 *Data analysis*

181 Two regressions were performed between the variables WCI and the AAR. The first regression was  
182 a linear least-square regression and the second was a non-linear regression fitting a logistic curve of  
183 the general form

$$184 \quad y = \frac{1}{1 + e^{-bx-c}}$$

185 where  $y$  is the fitted value of continuous probability of mangrove presence in the range 0–1,  $b$  is the  
186 slope parameter, and  $c$  is the estimated centre of transition. The two models were then examined for  
187 statistical significance and compared using Akaike information criterion (AIC). The AIC is a measure  
188 of the relative quality of statistical models for a given set of data. The lower the value of the AIC  
189 number, the better the model. Models were fitted using R (ver. 3.3.3, xxx; 6 March 2017). The  
190 resulting logistic curve was evaluated by calculating the point of inflection and the maximum rate of  
191 change. The basic differential equation underlying the logistic curve was assumed to be as follows:

$$192 \quad \frac{d(\text{WCI})}{d(\text{AAR})} = \alpha \text{WCI}(1 - \text{WCI})$$

193 where  $\alpha$  is a positive constant. By taking the second differential, it can be shown from this equation  
194 that the point of inflection must occur when WCI is 0.5. The value of AAR at the point of inflection  
195 and maximum rate of change of WCI with AAR can be found using this equation and the derived  
196 logistic equation.

#### 197 **Results**

198 There was a significant linear relationship between the WCI and AAR over 20–30 years (regression  
199 analysis:  $F_{203} = 72.31$ ,  $P < 0.001$ ,  $R^2 = 0.259$ ; linear equation:  $\text{WCI} = 0.00048 \times \text{AAR} - 0.16$ ; see Fig.  
200 3). The logistic regression of the WCI against the AAR over 20–30 years produced the following  
201 equation:

$$202 \quad \text{WCI} = \frac{1}{1 + e^{(0.00224\text{AAR} - 3.013)}} \quad (\text{see Fig. 3})$$

203 This equation gives the probability of the presence of mangrove. Both the coefficient of AAR and  
204 the intercept were significant at the  $<0.001$  level. The linear regression had an AIC value of 1953 and



205 the logistic regression yielded an AIC value of 243. Therefore, it can be concluded that the logistic  
206 regression yielded the superior model.

207 The point of inflection and the maximum rate of change of WCI with AAR were derived from the  
208 logistic equation, using Eqn 1 (see Methods). The point of inflection occurred at WCI = 0.5 and AAR  
209 = 1368 mm. The maximum rate of change of WCI with AAR was 5.5% of WCI for 100 mm of AAR.  
210 The range in which the probability shifts by more than 5% per 100 mm of rainfall was between an  
211 upper limit of probability of 0.65 and a lower limit of 0.39. It follows that in the range of 1086–1651  
212 mm of AAR, the probability of the presence of mangroves shifts positively by 5% per 100 mm of  
213 rainfall.

## 214 Discussion

### 215 *Tidal wetlands redefined*

216 Tidal wetlands along tropical and subtropical coastlines worldwide form a distinct ecological niche  
217 between the mean sea level and the highest tide level. This upper tidal niche, regularly inundated and  
218 washed by tidal waters and estuarine flows, is occupied by two well recognised vegetation types of  
219 mangroves and tidal saltmarsh with saltpans. Plants of these unique vegetation types share specialised  
220 physiological, morphological and anatomical traits for coping with salt and regular inundation by river  
221 and sea waters. This is in deference to their often quite different ancestral lineages and genetic make-  
222 up (Tomlinson 2016; Duke 2017). Such inherent differences explain in part why the various species of  
223 mangrove and saltmarsh plants respond differently to overall environmental drivers such as rainfall.  
224 Where rainfall conditions differ, these broad vegetation types appear to adjust their relative cover,  
225 while occupying the same combined area. Accordingly, estuarine sites with a higher rainfall had  
226 greater proportions of mangroves and, correspondingly, smaller areas of tidal saltmarsh and saltpan  
227 (see Fig. 1). In locations where rainfall levels were relatively low, the opposite was observed, with a  
228 reduced proportion of mangroves and a greater extent of saltmarsh and saltpan. Although it remains to  
229 be fully appreciated, the most plausible explanation is a dynamic process where changes in prevailing  
230 rainfall conditions lead to adjustments and changes in the respective amounts of dominant vegetation  
231 types. Accordingly, it is proposed that the relative extent of tidal-wetland plants is influenced by  
232 moisture availability, affecting the presence and growth of taller, high-biomass plants that capture  
233 light and out-compete their shorter co-inhabitants in circumstances when rainfall levels are sufficiently  
234 high.

235 Our current assessment re-examined the overall relationship between mangrove and saltmarsh  
236 vegetation, and AAR, as observed by Fosberg (1961). We found that the relationship was best  
237 described by a sigmoidal curve rather than a linear one. This sigmoidal relationship gave a measure of  
238 the probability of the presence of mangroves as a function of AAR. In so doing, the relationship has  
239 been usefully described and quantified as a model defined by proposed shifts in the ecotone between



240 mangrove and tidal saltmarsh–saltpan vegetation. The model specifically depicts a range of AAR  
241 between 1086 mm and 1651 mm, where the probability of the presence of mangroves increases by 5%  
242 per 100 mm of AAR.

### 243 *Geographic characterisation of the relationship with rainfall*

244 The relationship between rainfall and WCI can be used also to identify areas at risk of change in the  
245 composition of tidal-wetland vegetation. In Fig. 4, zones of predicted WCI values are shown for  
246 northern Australia, as derived from the 30-year AAR map (1976–2005). The time period corresponds  
247 with the majority of OzCoasts estuarine-mapping data taken from imagery collected up to 2005. It  
248 should also be noted that upland areas are shaded well inland because tidal wetlands are influenced by  
249 rainfall throughout their larger catchment areas, as well as on the downstream tidal-wetland areas  
250 themselves (Smith and Duke 1987; Duke et al. 1998). This tidal-wetland risk map can be used in the  
251 following two ways (see Fig. 4): (1) to compare predicted vegetative states with observed values at  
252 specific sites; and (2) to show tidal-wetland areas that are likely to have specific vegetative states,  
253 along with the underlying risks and vulnerabilities associated with each. It was evident that observed  
254 vegetative states at specific sites broadly matched the predicted WCI zones.

255 Furthermore, on the basis of the relationship between the AAR and WCI (Fig. 4), notable changes  
256 may be expected along the tidal-wetland ecotone in response to changes in AAR. These effects are  
257 likely to manifest as an ecotone shift, with either mangrove expansion or mangrove loss being  
258 expected under the following scenarios: (1) increases of ~100 mm in AAR in zones with predicted  
259 WCI values between 40 and 80% (mapped as pink and green zones) might result in gains up to 10% in  
260 the WCI, equating to significant mangrove expansion; or (2) declines of ~100 mm in AAR in zones  
261 with predicted WCI values between 20 and 40% (mapped as pink and yellow zones) might result in  
262 losses up to 10% in the WCI, equating to severe mangrove dieback. The likelihood of such changes is  
263 expected to differ according to these zones where, for example, outside the zones (i.e. in regions  
264 receiving <500 mm or >2000 mm AAR) changes of 200–1000 mm in AAR would have a limited  
265 effect on the WCI, and the proportional extent of mangrove cover would, therefore, remain relatively  
266 constant. The position of an estuarine system within a zone exhibiting a predicted WCI value between  
267 20 and 80% (mapped as yellow, pink and green in Fig. 4) is most pertinent for estuarine systems that  
268 have been subject to either long-term increases or long-term declines in the AAR over the past half  
269 century. In Fig. 4, the zone of pink shading with red-dashed line indicates those regions falling near  
270 the ‘inflection point’ in vegetative status, making them most at risk of significant impacts from  
271 relatively small changes in the AAR ( $\pm 100$  mm, see below). As such, the ‘inflection point line’ or  
272 zone identifies both areas most at risk, and the dividing line between mangrove-dominated and tidal  
273 saltmarsh–saltpan-dominated wetlands. It is of interest that this zone of relatively unstable conditions  
274 includes most coastal areas of northern Australia.

275 **Further deductions from the relationship**

276 Because the spatial evidence presented portrays a range of wetland-cover (WCI) states correlated  
277 with rainfall, it is reasonable to assume that these states could have been achieved only if past  
278 vegetative states had responded to ambient rainfall conditions. And, as rainfall conditions changed, the  
279 corollary is that vegetative states need to have responded for the correlation to hold. Therefore, it can  
280 be deduced from these spatial data that changes between mangrove-dominated and saltmarsh–saltpan-  
281 dominated vegetative states appear to exist and function in a state of hysteresis. There is good reason  
282 to believe that vegetative condition depends on the on-going condition of climate factors, in particular  
283 rainfall. A range of possible states and intermediate conditions are portrayed in the conceptual model  
284 (Fig. 5). Trends towards alternate extreme states depend on both spatial and temporal aspects of  
285 climate and levels of precipitation. In accordance with such circumstances, the model schematic  
286 describes a range of condition states defined by the WCI, where there might be expansion or  
287 contraction of mangroves in regions within their natural latitudinal range. The model depicts the  
288 extreme range of hypothetical states, with trends between possibly increasing wet- or dry-condition  
289 states. To fully understand how these habitat responses in vegetation cover take place, case studies of  
290 historical change are needed to exemplify how temporal changes in vegetation cover match past  
291 changes in climate.

292 **Overall conclusions**

293 Our findings affirmed Fosberg's (1961) assertion of a significant and meaningful relationship  
294 between vegetative cover of tidal wetlands and rainfall of tropical and subtropical areas. In addition,  
295 the relationship describes the strong likelihood of a functionally dynamic ecosystem alternating  
296 between high- and low-biomass states determined by longer-term trends in rainfall and climate. This  
297 observation is likely to have wider implications when considering associated and dependent biota, as  
298 well as the overall productive capacity of these habitats and their delivery of ecosystem services. The  
299 model developed helps explain how the relative cover of mangrove and saltmarsh vegetation of tidal-  
300 wetland systems are demonstrably predictable, provided local, direct human pressures are minimal.

301 Changes are also likely manifest as longer term, decadal-scale shifts in vegetation ecotones between  
302 high-biomass and low-biomass tidal-wetland communities. In each case, the relationship identified  
303 provides a means to extrapolate levels of longer-term rainfall from current observations of extent of  
304 tidal-wetland plant assemblages. Although case studies of change to specific estuarine systems are  
305 needed (cs. Nicholls *et al.* 1997), it will be important to consider whether the model described in the  
306 present treatment can help explain each instance of loss or gain within tidal-wetland plant  
307 assemblages. Future mapping of tidal-wetland areas must at least include concurrent measures of both  
308 mangrove and tidal saltmarsh–saltpan communities, as well as the clear attribution of the dates of  
309 source imagery.



310 We identified a marked ecotone at the interface between distinctly high- and low-biomass  
311 vegetation types. This dynamic interface is likely to manifest further as the location for encroachment–  
312 expansion or dieback–contraction to match corresponding increases or decreases in longer term  
313 rainfall. Future case studies are needed to record how, where and when mangrove vegetation might  
314 change as rainfall conditions change. They are also needed for an evaluation of the role of moisture  
315 stress to help explain each significant instance of loss or gain within tidal-wetland assemblages.

### 316 Conflicts of interest

317 The authors declare that they have no conflicts of interest.

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468 **Table 1. Summary of available data for an estuarine site of minimal disturbance grouped**  
 469 **according to average catchment rainfall, areas of key tidal-wetland vegetation types (mangrove,**  
 470 **saltmarsh–saltpan) and estimates of wetland cover index (see text) for three Australian states**

471 WA, Western Australia; NT, Northern Territory; Qld, Queensland. Source: Australian Government

472 Geosciences Australia (2017)

Data feature	WA	NT	Qld	Total
Mangrove-species diversity*	19	32	45	46
Total in all of national database				
Estuarine systems in database	227	151	322	770
Area of mangrove (km <sup>2</sup> )	1985	3289	3215	8560
Area of tidal saltmarsh–saltpan (km <sup>2</sup> )	5137	4925	5253	15442
Wetland cover index (%)	27.9	40	38	35.7
Estuarine systems in study				
Pristine- and medium-area systems	35	76	94	205
Area of mangrove (km <sup>2</sup> )	485	1255	1095	2834
Area of tidal saltmarsh–saltpan (km <sup>2</sup> )	494	1425	1800	3720
Total tidal-wetland area (km <sup>2</sup> )	979	2680	2895	6554
Mean annual rainfall (mm)	911	1339	1278	1248
Wetland cover index (%)	49.5	46.8	37.8	43.2

473 **Fig. 1.** Examples of tidal-wetland vegetation (mangrove plus tidal saltmarsh–saltpan, as high- and low-biomass  
 474 states respectively), showing variations in percentage mangrove cover in relation to mean annual rainfall (Duke  
 475 *et al.* 2003) for four sites in north-eastern Australia: (a) ~96% mangrove in Mutchero Inlet south of Cairns at  
 476 ~3600 mm rainfall (17°14′01.75″S, 145°57′03.00″E); (b) ~75% mangrove in Cattle Creek near Ingham at ~1800



477 mm rainfall (18°52'40.91"S, 146°16'59.95"E); (c) ~47% mangrove in Alligator Creek south of Townsville at  
478 ~1032 mm rainfall (19°18'34.68"S, 146°55'56.27"E); and (d) ~28% mangrove on Balaclava Island off the mouth  
479 of the Fitzroy River, at ~700 mm rainfall (23°34'31.35"S, 150°56'00.07"E).

480 **Fig. 2.** Mangrove cover in relation to (a) rainfall and (b) Australian state or territory, as quantified by the  
481 wetland cover index (WCI, present study) and the BS index (Bucher and Saenger 1994). WCI expresses  
482 estuarine mangrove cover as a proportion of the sum of cover of intertidal habitats (mangroves, saltmarshes and  
483 salt pans). BS quantified mangrove cover as a proportion of intertidal as well as subtidal habitats (mangroves,  
484 saltmarsh/clay pan, seagrass, lower intertidal flats and open water). Indices were calculated from the dataset used  
485 by Bucher and Saenger (1994).

486 **Fig. 3.** The relationship between tidal-wetland vegetation cover and average annual rainfall in estuarine sites  
487 across northern tropical Australia. The 205 estuarine locations covered a range of climatic conditions from semi-  
488 arid to very wet conditions. Both linear and logistic regressions were significant. Based chiefly on the logistic  
489 curve, the pink-shaded selection of sites, bordered by yellow (more arid) and green (wetter) groupings, depict an  
490 inflection point (1000–1500 mm average annual rainfall) between saltmarsh–saltpan-dominated (low-biomass)  
491 systems and mangrove-dominated (high-biomass) systems.

492 **Fig. 4.** A map of northern Australia depicting areas at risk of change (cs. mangrove loss vs mangrove  
493 expansion) as a result of changes in longer-term trends in rainfall. Model parameters were based on the  
494 comparison of observed and predicted levels of dominance of tidal-wetland vegetation types before 2006, which  
495 were derived from area data sourced from the OzCoasts online database (Australian Government Geosciences  
496 Australia, 2017; see Table 1), and mean annual-rainfall data (Australian Government Bureau of Meteorology  
497 2017; central inset). Coastal areas occupied by mangroves are shown with dark blue shading (modified from  
498 Duke 2006). Landward shaded zones show five levels of predicted vegetative dominance (see Methods and Fig.  
499 3) for amounts of mean annual rainfall in respective coastal and estuarine catchment areas. The pink zone and  
500 red dashed line represent areas of greatest likelihood of change in vegetation cover, where annual rainfall had  
501 trended towards either wetter or drier conditions (see Fig. 3, 5). Individual site records in the OzCoasts database  
502 were plotted as coloured-circle symbols at respective coastal locations, each according to specific levels of WCI.  
503 The main legend (lower centre) describes the five mangrove-dominance groupings, rainfall conditions and  
504 summary parameters of predicted levels of risk.

505 **Fig. 5.** Proposed model of ecotone shift in mangrove tidal wetlands. The cycle has opposing extreme states in  
506 wet and dry climates respectively. Intermediate states are normal, with marked ecotones between mangrove and  
507 saltmarsh–saltpan vegetation types present. These observations describe the possible hysteresis pathway where  
508 the ecotone between mangroves and saltmarsh–saltpan shifts in predictable ways with longer-term trends in  
509 rainfall.